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# Gravity wave propagation in inhomogeneous media : wave scattering and interference process

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# Vincent REY, hydrodynamics RESEARCH ACTIVITIES

- Wave propagation and transformations in the coastal zone
- Sediment transport and morphodynamics

## Applications

Environmental nearshore dynamics



Ocean and coastal engineering



# CONTENTS

## → Some generalities on gravity waves

- Conservation equations and velocity potential
- Wave properties

## → Resonant interactions (2D cases) : wave reflection and pressure oscillations

- Standing waves : basin oscillations or « seiching »
- Interference processes : application to nearshore structures
- Pressure oscillation : application to energy power device

## → Refraction – diffraction (3D cases) : energy focusing and resonant interactions

- Wave celerity
- Wave focusing
- Resonant interactions : Bragg resonance
- Application to energy power device : Oscillating Water Column

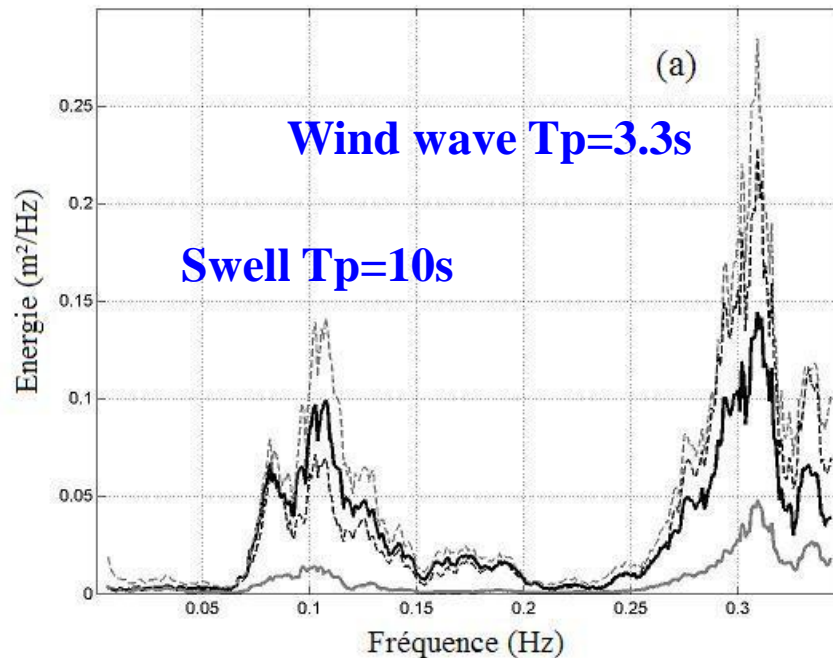
# Water waves : sea states and wave models



Wind wave (+ swell)



Swell (quasi-sinusoidal shape)



Sea states:  
Waves characteristics :  $H_s$ ,  $T_{peak}$ , etc..



Wave models

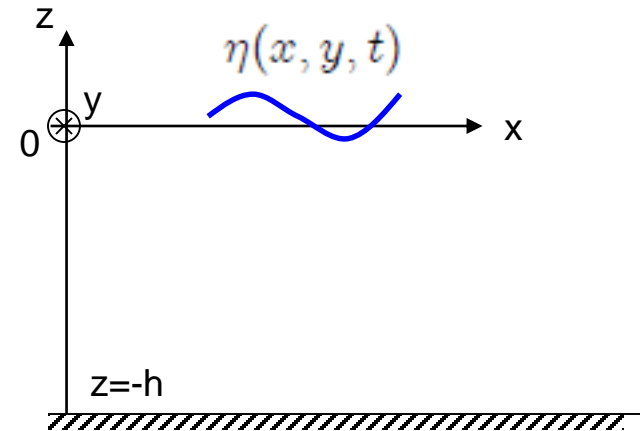
- Monochromatic waves (phase-resolving models)
- Spectral waves (phase averaged models, (based on the energy flux conservation))

# MONOCHROMATIC WAVES : STOKES THEORIES

## Hypothesis

- Inviscid fluid : Euler equations
- Irrotational motion

→ Potential flow  $\Phi(x, y, z, t)$  and  $\vec{v} = \nabla\Phi$



## Conditions for the potential $\Phi(x, y, z, t)$

- Laplace's equation ( $\eta > z > -h$ ):  $\nabla^2\Phi = 0$

- Free surface conditions ( $z = \eta$ ) :

- Kinematic :

$$\frac{\partial\eta}{\partial t} + \frac{\partial\Phi}{\partial x} \frac{\partial\eta}{\partial x} - \frac{\partial\Phi}{\partial z} = 0$$

- Dynamic (Bernoulli) :


$$g\eta + \frac{1}{2}\vec{v}^2 + \frac{\partial\Phi}{\partial t} = 0$$

- Bottom impermeability ( $z = -h$ ) :

$$\frac{\partial\Phi}{\partial z} = 0$$

using  $\vec{v} \cdot \vec{\nabla} \left( \frac{\partial \Phi}{\partial t} \right) = \frac{1}{2} \frac{\partial \vec{v}^2}{\partial t}$

Total differentiation of the dynamic free surface equation and choice  $p=0$  for the surface pressure lead to the free surface condition ( $z=\eta$ ) :

 
$$g \frac{\partial \Phi}{\partial z} + \frac{\partial \vec{v}^2}{\partial t} + \frac{\partial^2 \Phi}{\partial t^2} + \frac{1}{2} \vec{v} \cdot \vec{\nabla} (\vec{v}^2) = 0$$

## REMARKS

- The free surface condition is **NON-LINEAR**
- Expression is given at  $z=\eta$ , which depends on both time and space



Approximations:

- Linear : 1st order Stokes wave or Airy wave
- Perturbation methods : Stokes waves at higher orders

NON-LINEAR solutions : Asymptotic expansion of  $\eta$  and  $\Phi$

$$\begin{cases} \eta = \varepsilon\eta_1 + \varepsilon^2\eta_2 + \varepsilon^3\eta_3 + O(\varepsilon^4) \\ \Phi = \varepsilon\Phi_1 + \varepsilon^2\Phi_2 + \varepsilon^3\Phi_3 + O(\varepsilon^4) \end{cases}$$

Free surface condition for  $\Phi$  and Bernoulli expansion for  $\eta$  expressed at  $z=0$  using Taylor expansion:

$$f(x, \eta, t) = f(x, 0, t) + \eta \left[ \frac{\partial f}{\partial z} \right]_{z=0} + \frac{\eta^2}{2} \left[ \frac{\partial^2 f}{\partial z^2} \right]_{z=0} + \dots$$

Free surface (combined kinematic and dynamic (Bernoulli) conditions:

$$O(\varepsilon): \frac{\partial^2 \Phi_1}{\partial t^2} + g \frac{\partial \Phi_1}{\partial z} = 0$$

$$O(\varepsilon^2): \frac{\partial^2 \Phi_2}{\partial t^2} + g \frac{\partial \Phi_2}{\partial z} = -\eta_1 \left( \frac{\partial^3 \Phi_1}{\partial t^2 \partial z} + g \frac{\partial^2 \Phi_1}{\partial z^2} \right) - \frac{\partial}{\partial t} \left[ \left( \frac{\partial \Phi_1}{\partial x} \right)^2 + \left( \frac{\partial \Phi_1}{\partial z} \right)^2 \right]$$

Bernoulli condition :

$$O(\varepsilon): -g\eta_1 = \frac{\partial \Phi_1}{\partial t}$$

$$O(\varepsilon^2): -g\eta_2 = \frac{\partial \Phi_2}{\partial t} + \eta_1 \frac{\partial^2 \Phi_1}{\partial z \partial t} + \frac{1}{2} \left[ \left( \frac{\partial \Phi_1}{\partial x} \right)^2 + \left( \frac{\partial \Phi_1}{\partial z} \right)^2 \right]$$

## 1st Order STOKES Wave : Airy wave

Wave steepness  $ka \ll 1$ :

Free surface boundary condition ( $z=0$ ) : 
$$g \frac{\partial \Phi}{\partial z} + \frac{\partial^2 \Phi}{\partial t^2} = 0$$

For a free surface deformation of the form  $\eta(x, t) = a \sin(\omega t - kx)$

The potential  $\Phi(x, z, t)$  is given by

$$\Phi(x, z, t) = \frac{a\omega}{k} \frac{\cosh[k(z+h)]}{\sinh(kh)} \cos(\omega t - kx)$$

Elliptic  
trajectories

$$X = a \frac{\cosh[k(z+h)]}{\sinh(kh)} \cos(\omega t - kx) = X_0 \cos(\omega t - kx)$$

$$Z = a \frac{\sinh[k(z+h)]}{\sinh(kh)} \sin(\omega t - kx) = Z_0 \sin(\omega t - kx)$$

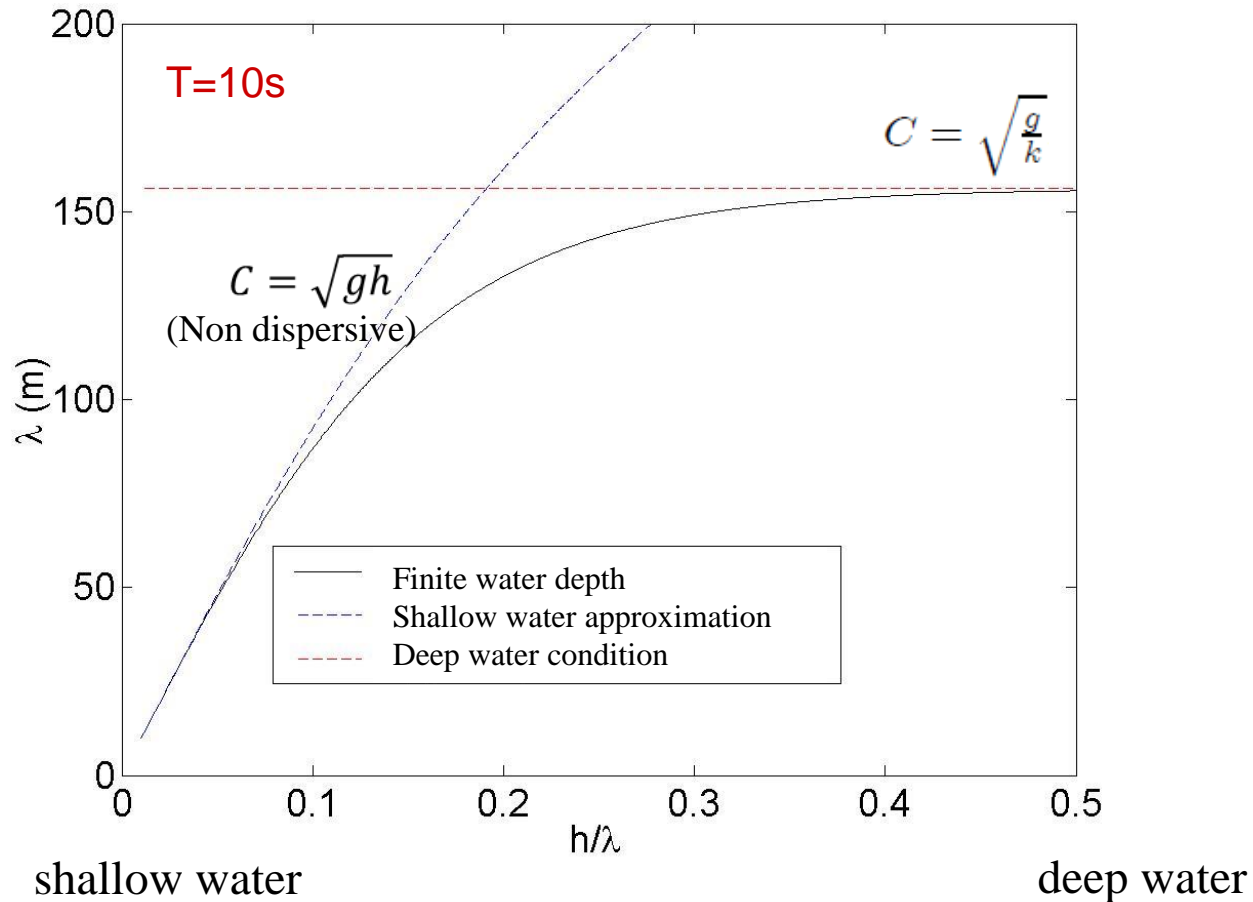
$$\left[ \frac{X}{X_0} \right]^2 + \left[ \frac{Z}{Z_0} \right]^2 = 1$$



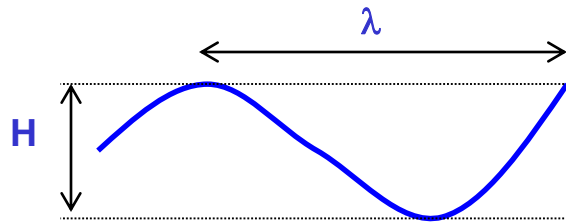
The wavenumber  $k$  verifies the dispersion relation  $\omega^2 = gk \tanh(kh)$

The wave celerity (phase celerity) is given by  $C = \frac{\omega}{k} = \sqrt{\frac{g}{k} \tanh(kh)}$

➡ The wave propagation is **dispersive**



# CHARACTERISTIC DATA and CLASSIFICATIONS



$T \approx 3-6s$  wind waves

$T \approx 8-14s$  Swell

Deep water

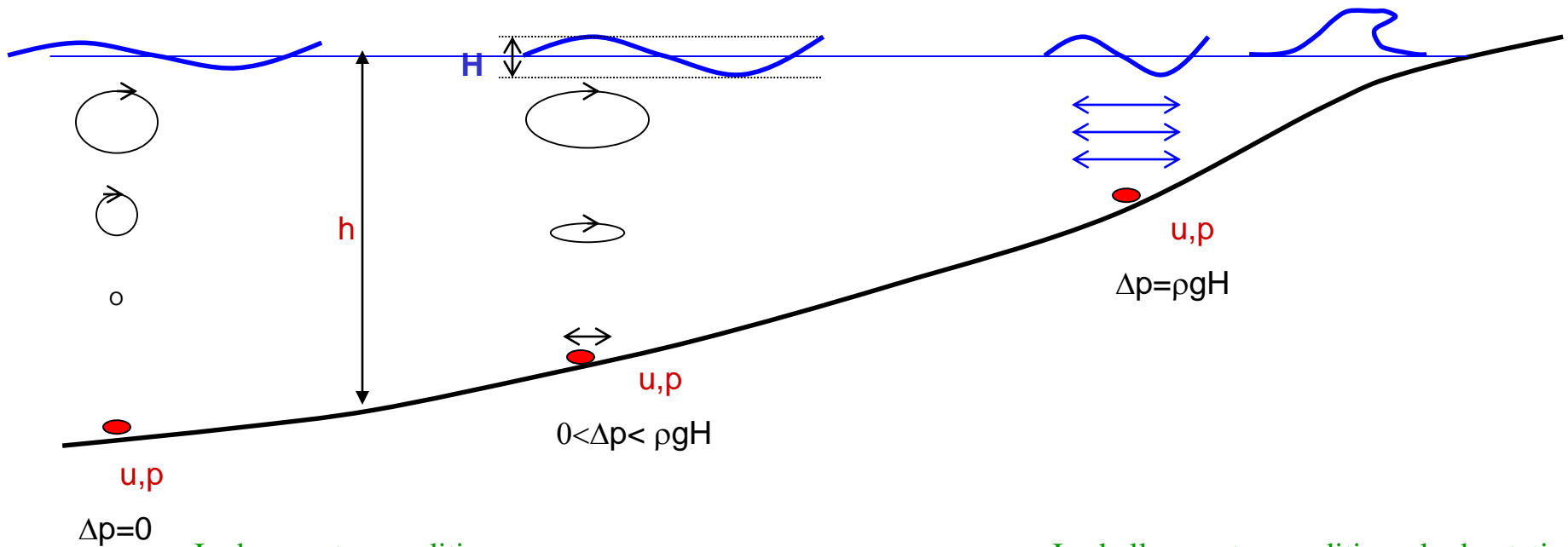
Intermediate water depth

Shallow water

$h > \lambda/2$

$\lambda/2 > h > \lambda/10$

$h < \lambda/10$



In deep water conditions, no wave impact near the bottom

In shallow water conditions, hydrostatic conditions (in the linear approximation)

## Wave energy :

➔ **Potential energy** by unit volume:  $d^3 E_p = \rho g z dx dy dz$

**Potential energy** by unit wavelength  $\lambda$  and width  $dy$  for a monochromatic wave propagating along the x-axis :

$$E_p = \int_0^\lambda \rho g \frac{\eta^2(x)}{2} dx$$

**Airy wave:**

$$E_p = \frac{\rho g a^2}{4} \lambda$$

➔ **Kinetic energy** by unit wavelength  $\lambda$  and width  $dy$  :

$$E_c = \frac{1}{2} \int_{x=0}^\lambda \int_{z=-h}^\eta \rho \vec{v}^2 dx dz$$

$$E_c = \frac{\rho g a^2}{4} \lambda$$

➔ **Total energy** by unit length  $dx$  and width  $dy$  :

$$E = \frac{\rho g a^2}{2}$$

➔ **Mean energy flux** across the  $y_0z$  plane, normal to the direction of propagation:

$$E_t = \frac{1}{T} \int_t^{t+T} \int_{x=-h}^\eta p \vec{v} \cdot \vec{n} dz dt \quad \text{with} \quad p = -\rho \frac{\partial \Phi}{\partial t} - \frac{1}{2} v^2$$

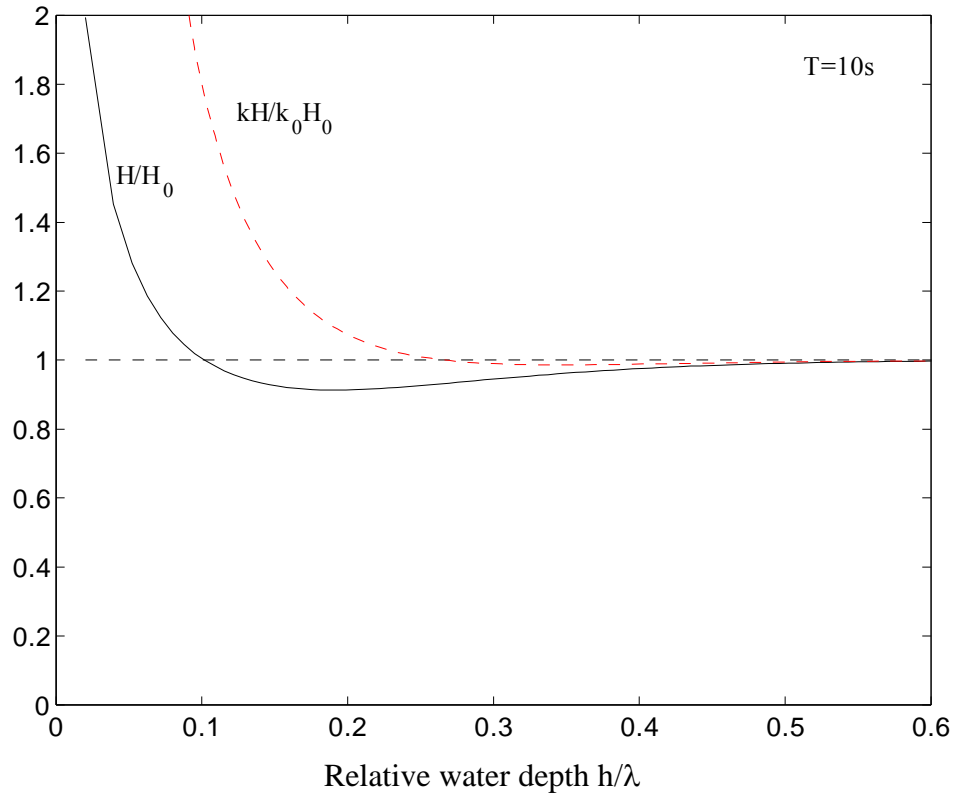
**Airy wave :**

$$E_t = \frac{\rho g a^2}{4} C \left[ 1 + \frac{2kh}{\sinh 2kh} \right] = \frac{\rho g a^2}{2} C_g = EC_g \quad \text{with} \quad C_g = \frac{\partial \omega}{\partial k} = \frac{1}{2} C \left[ 1 + \frac{2kh}{\sinh 2kh} \right]$$

Wave energy is transported at the group velocity

Wave propagation on a gentle slope, the « shoaling »  
 Energy flux conservation for a progressive wave:

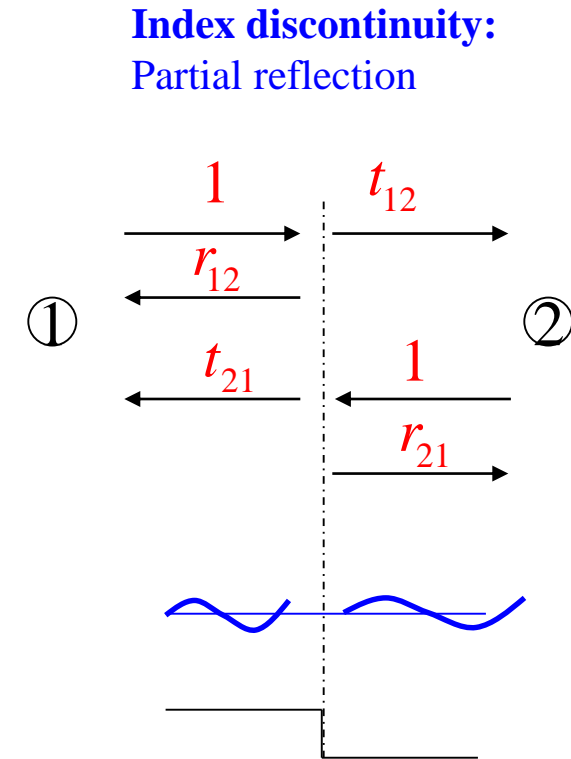
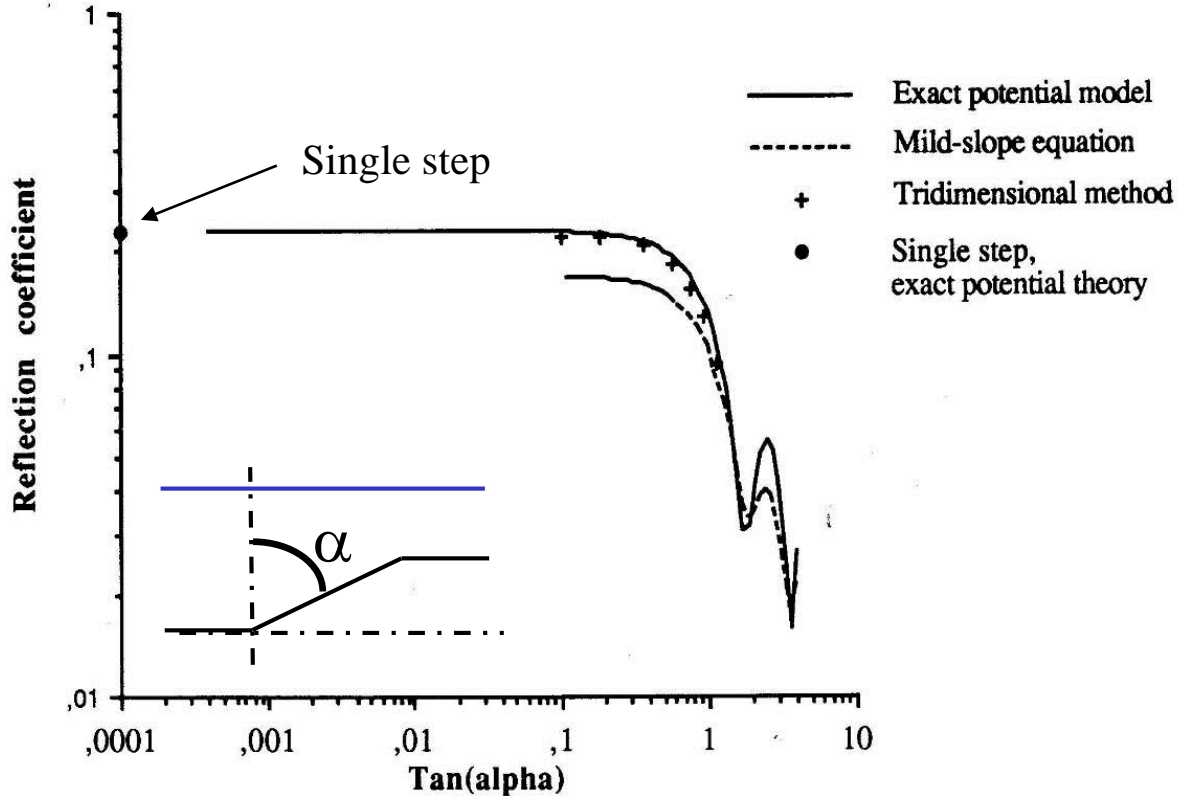
$$\frac{\partial E_t}{\partial x} = 0 \quad \longrightarrow \quad \left(\frac{a}{a_0}\right)^2 = \left(\frac{H}{H_0}\right)^2 = \frac{C_{g0}}{C_g} = \frac{1}{\tanh kh + \frac{kh}{\cosh^2 kh}}$$



Wave steepness  $kH$  rapidly increases when  $h$  decreases

→ Non linear effects

# Effect of a bed slope on the wave reflection

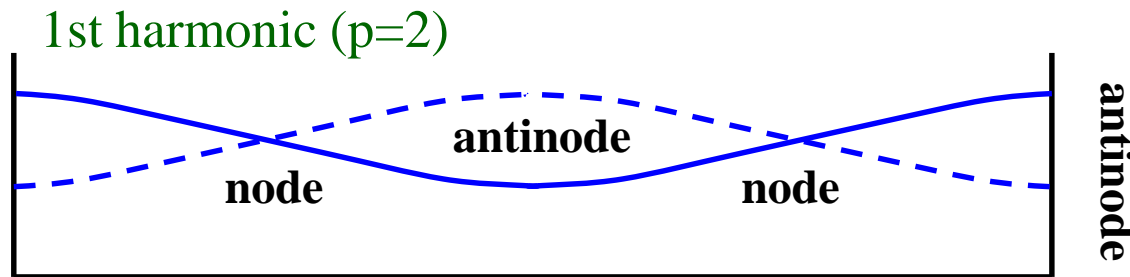
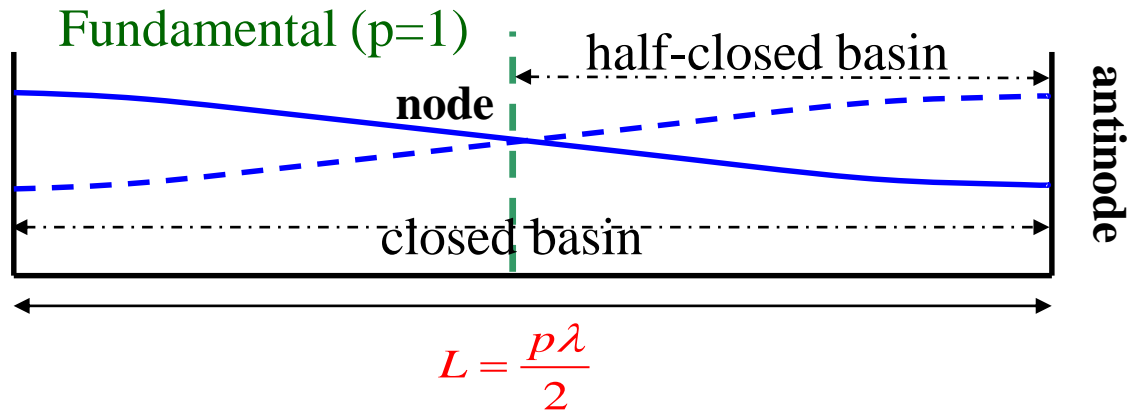


Gentle slopes → weak reflection (< 1% in terms of energy)

Bottom step → reflection cannot be neglected

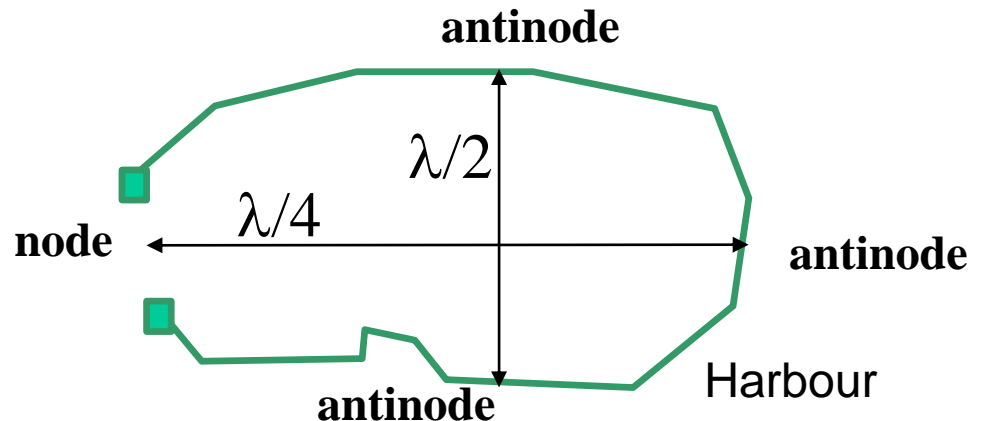
# Resonant interactions (2D cases)

Standing waves : basin oscillations or « seiching »



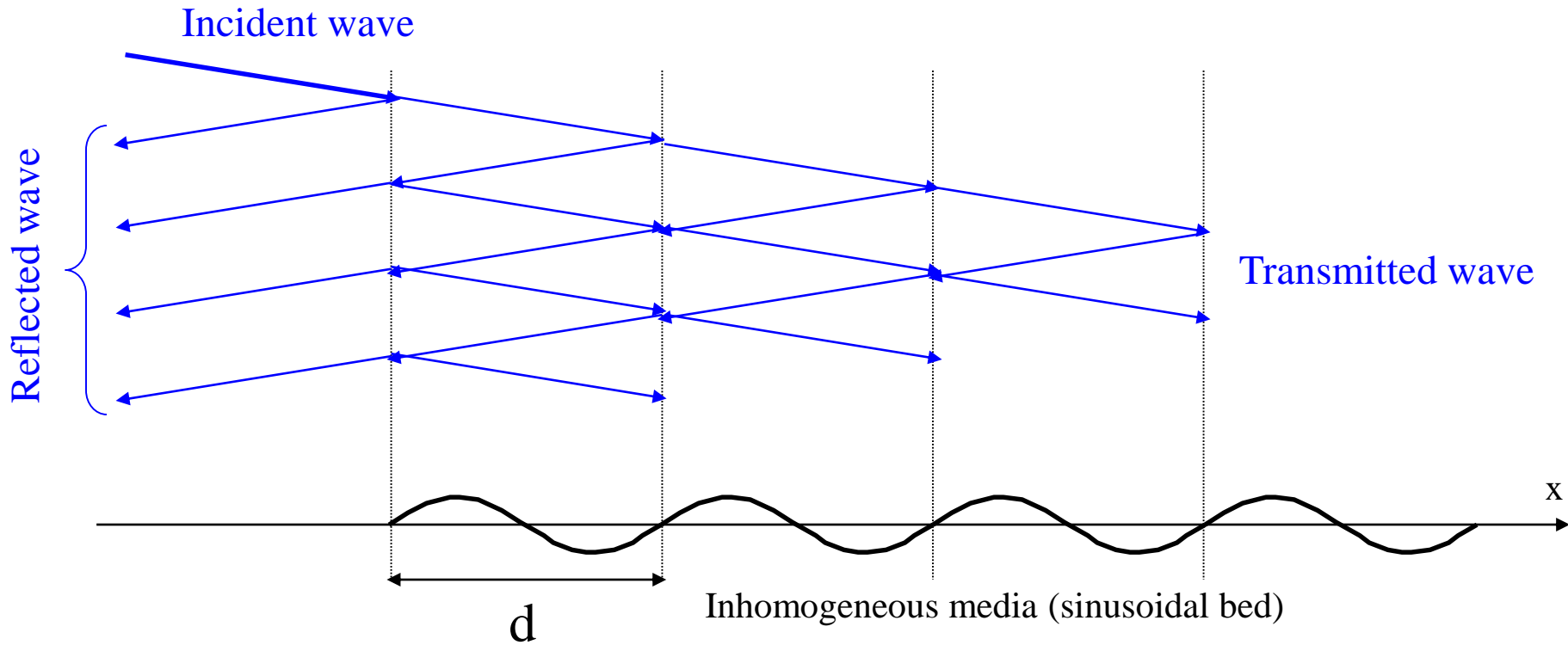
→ Harbour oscillations

Periods of seiching :  
Several to tens minutes



# Partially standing waves : Interference processes

## Interferences : physical approach



Phase lag after one return path:  $\Delta\varphi = 2\frac{2\pi}{\lambda} d$

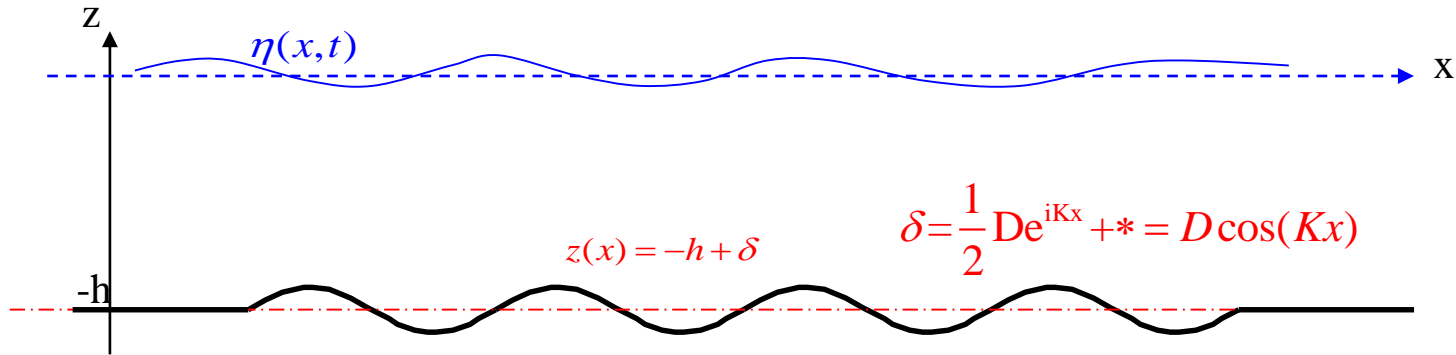
Phase matching  $\longrightarrow$  Reflection « allowed »

Opposition of phases  $\longrightarrow$  Reflection « forbidden »

$\longrightarrow$  Interference process

$\longrightarrow$  Maximum of reflection for  $\lambda=2d$  (Bragg Resonance)

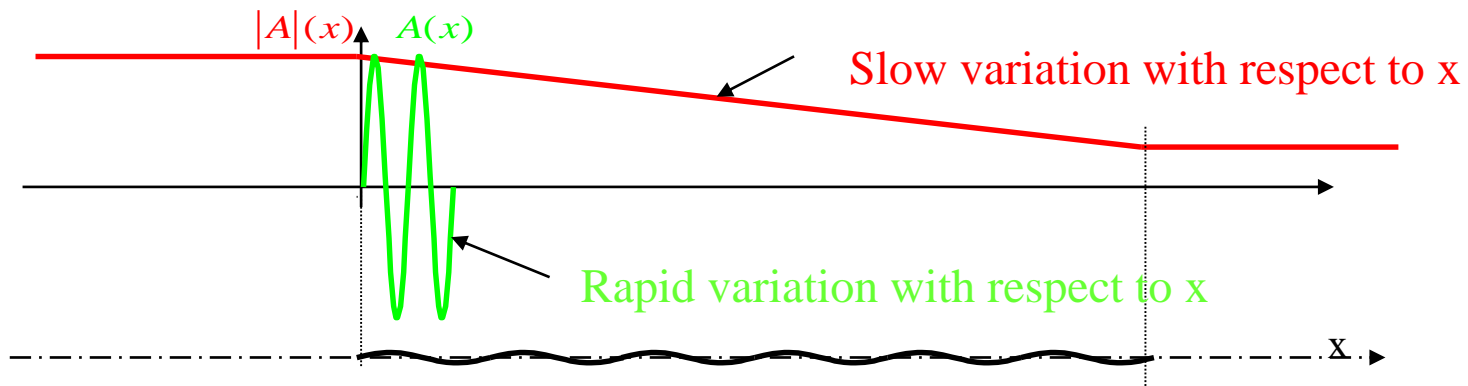
## Reflection forced by the bottom boundary condition



Bottom boundary condition  $\frac{\partial \Phi}{\partial z} = -\frac{\partial \delta}{\partial x} \frac{\partial \Phi}{\partial x}$  for  $z = -h + \varepsilon \delta$

Taylor expansion ( $z = -h$ ):  $\frac{\partial \Phi}{\partial z} = \frac{\partial}{\partial x} \left[ \delta \frac{\partial \Phi}{\partial x} \right] + \dots$

If  $K = 2k$ :  $\frac{\partial}{\partial x} \left[ \delta \frac{\partial \Phi}{\partial x} \right] \propto \cos(\omega t + kx) \quad \rightarrow \text{reflected wave}$



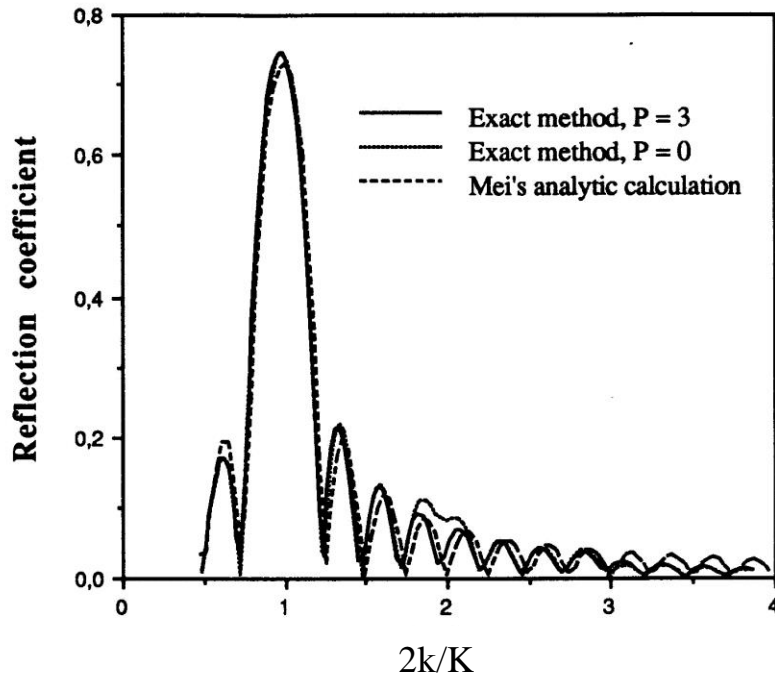
$\rightarrow$  **Multi-scale expansion method:** slow variables  $x_1 = \varepsilon x$  and  $t_1 = \varepsilon t$

$$\Phi(x, z, t) = \varepsilon \Phi_1(x, z, t, x_1, t_1) + \varepsilon^2 \Phi_2(x, z, t, x_1, t_1) + \dots$$



# Reflection forced by the bottom boundary condition

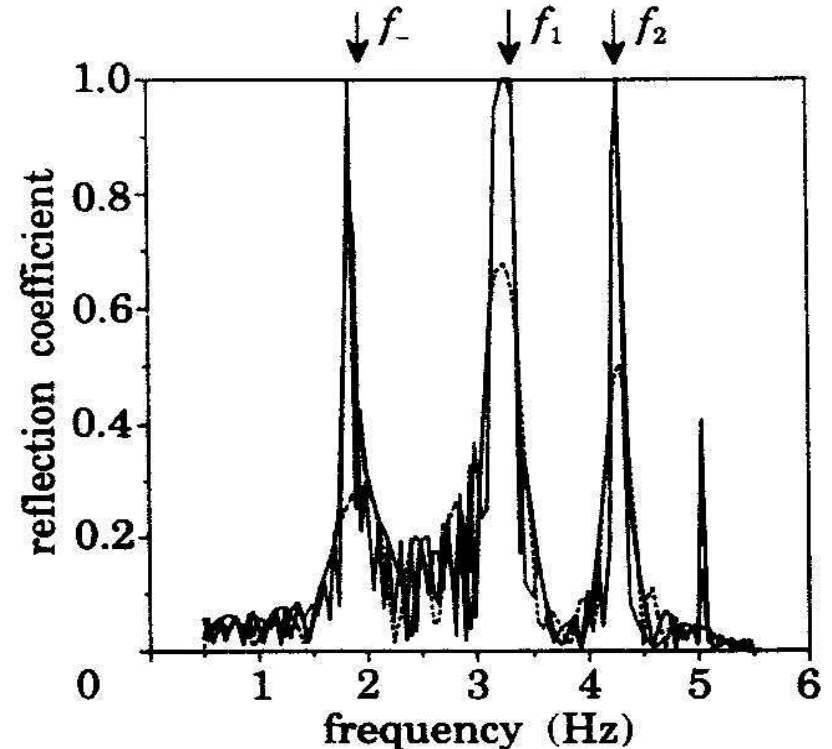
## Sinusoidal bed



Bragg resonance for  $K=2k$  :

- ➔ Strong reflection at the Bragg conditions
- Oscillations on both parts due to the finite length of the bed

## Doubly sinusoidal bed



1st order and subharmonic (2<sup>nd</sup> order) Bragg resonance

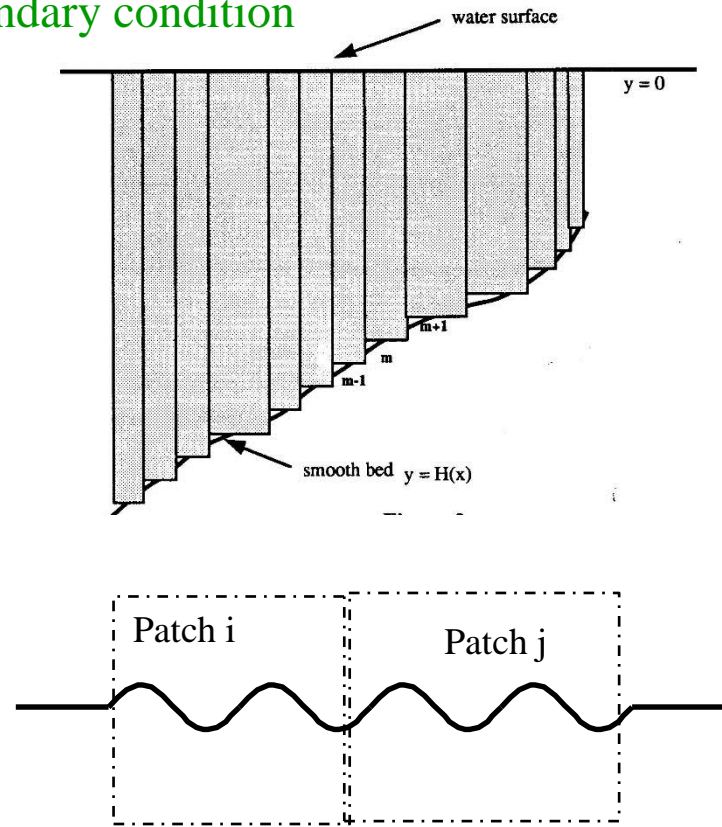
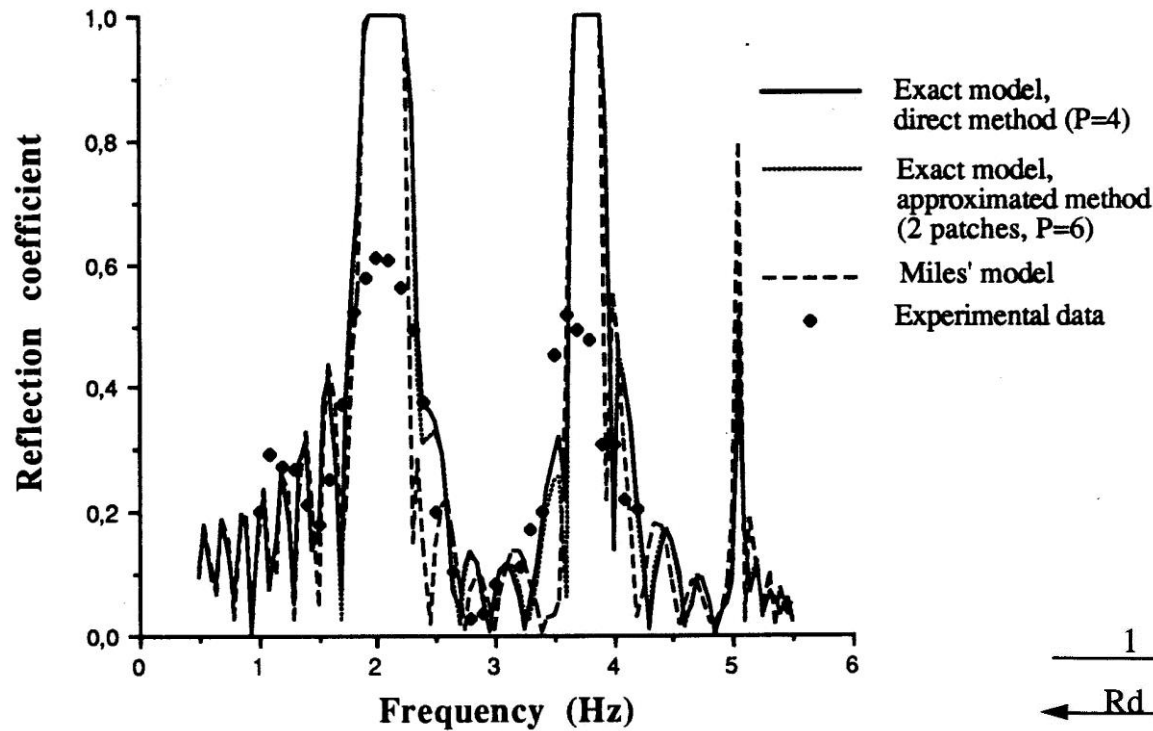
- ➔ High interaction at low frequency even if « 2<sup>nd</sup> order » Bragg
- ➔ Reflection tends to 1 for increasing bed length

# Reflection forced by the bottom boundary condition

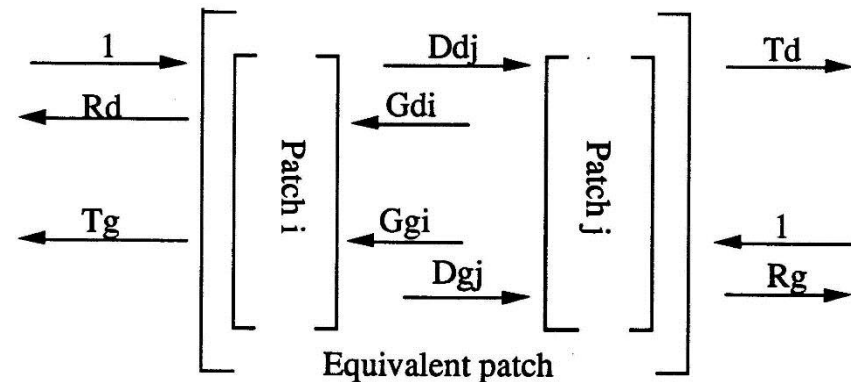
Doubly sinusoidal bed

Interference process :

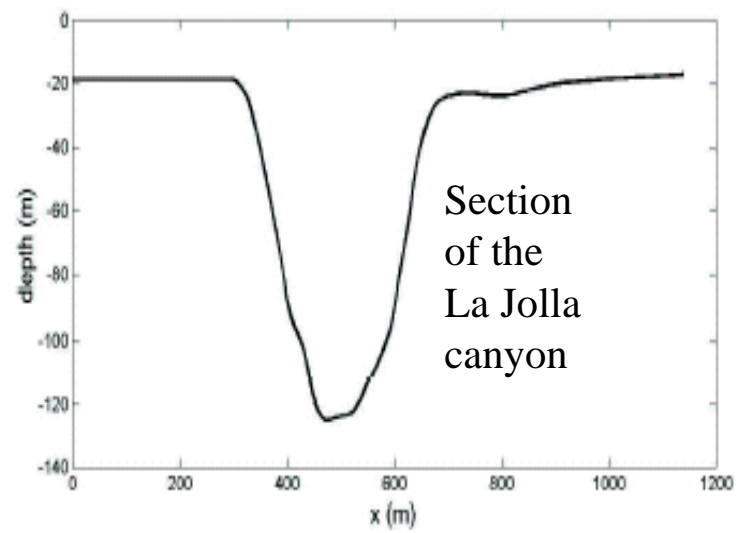
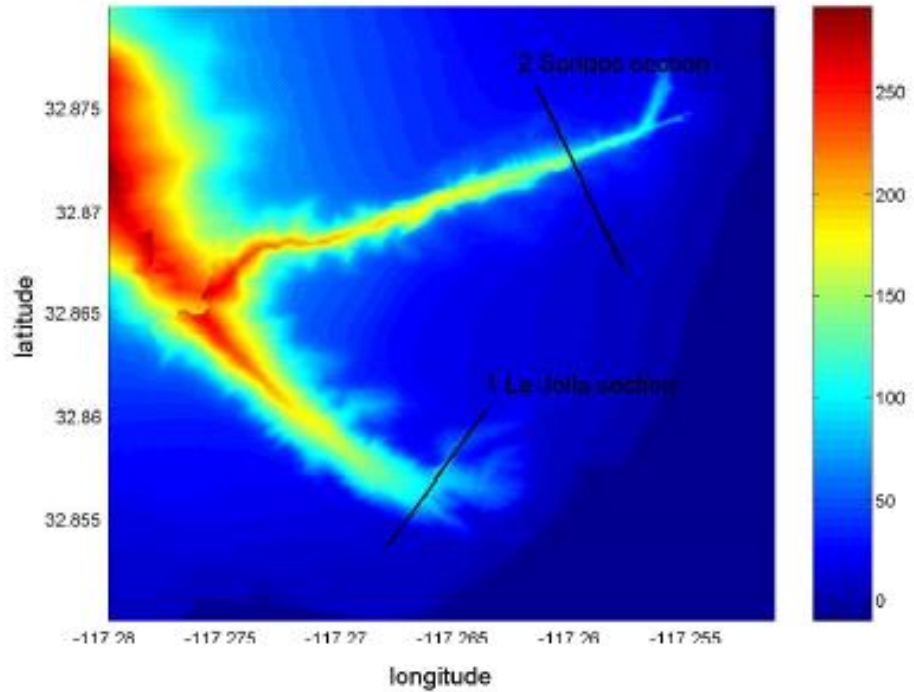
Analogy with the PF interferometer



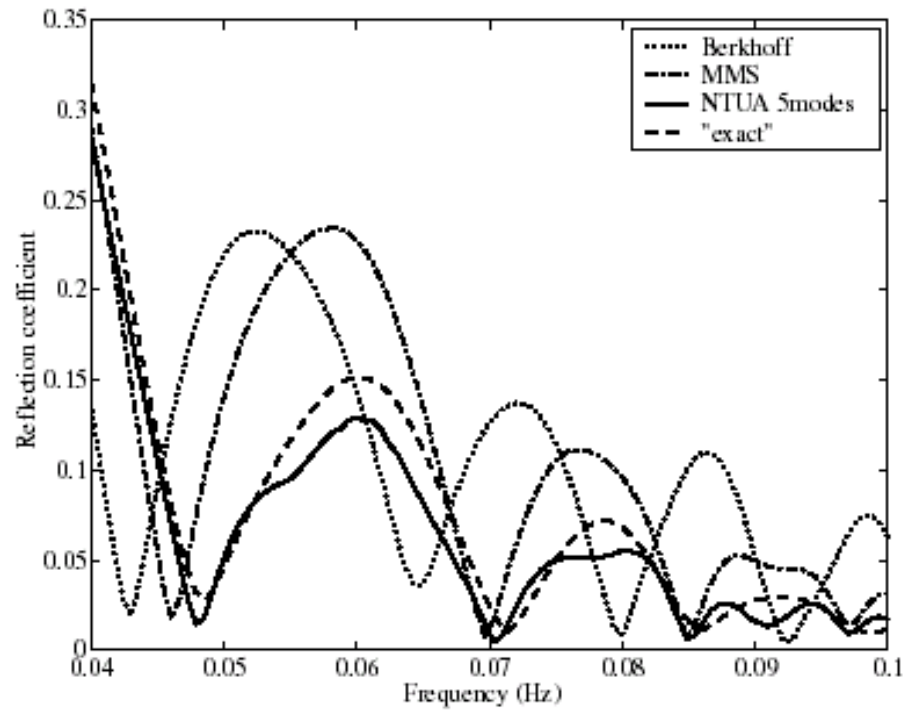
interference process



# Reflection in the presence of canyon : the La Jolla canyon (california, USA)

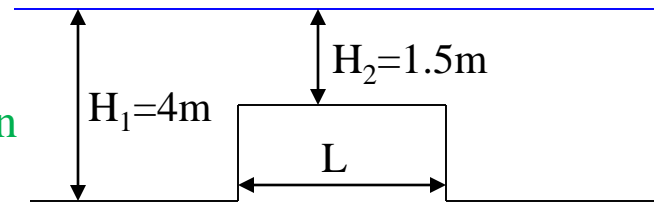


## Reflection coefficient for a normally incident wave

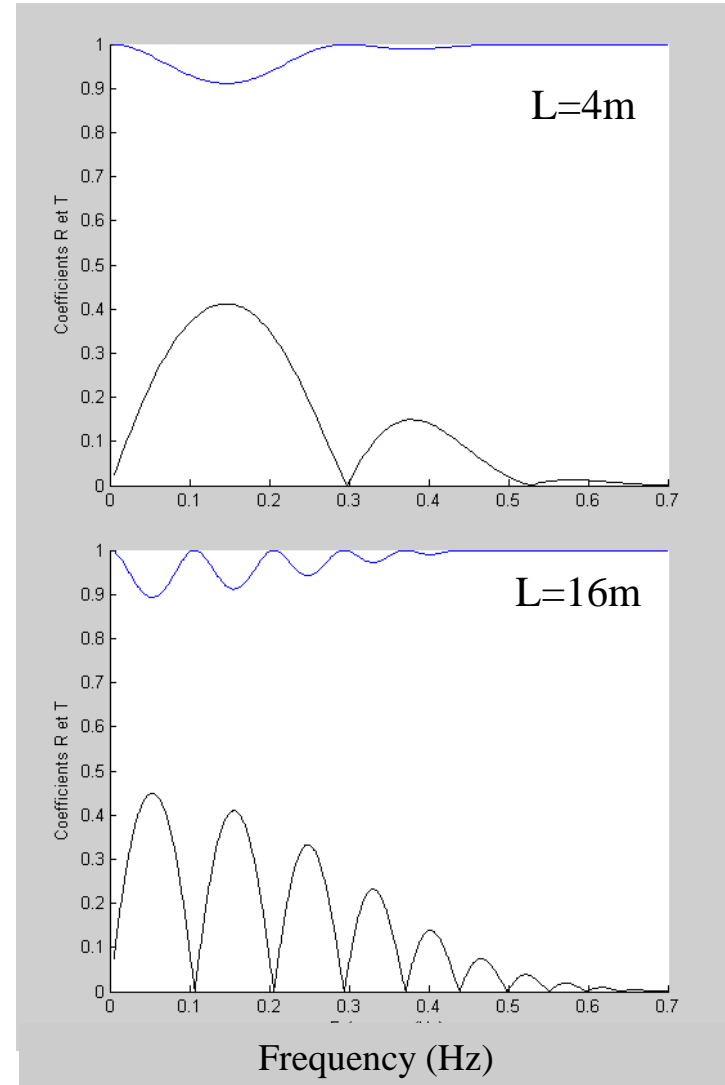
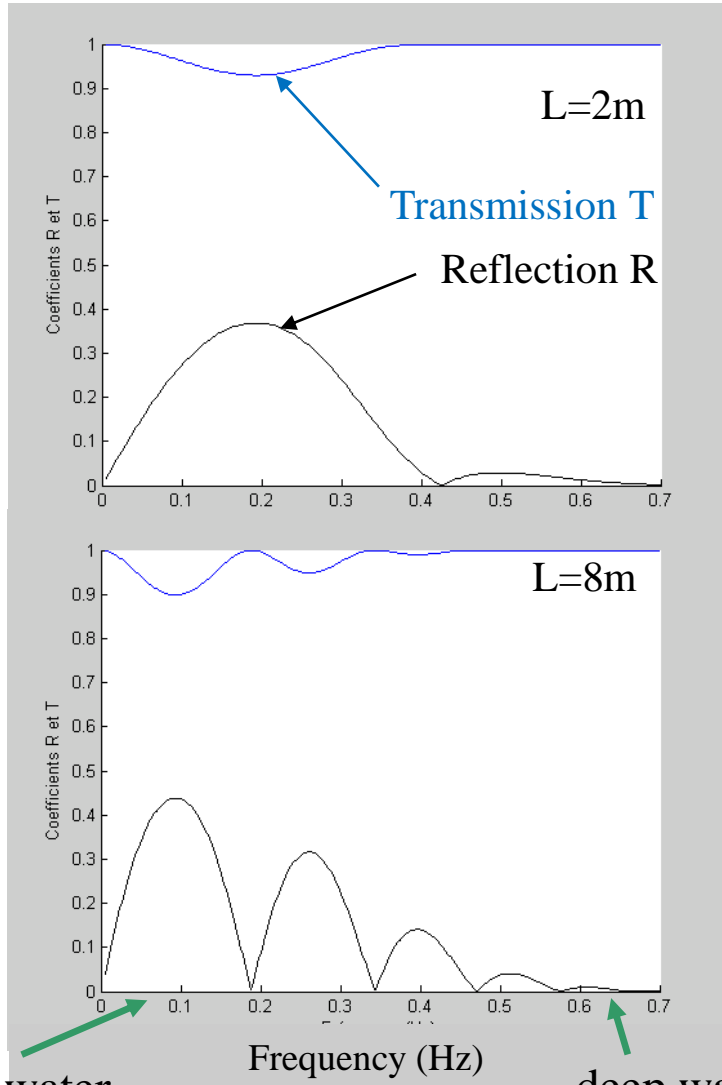


# Application to nearshore structures:

## Rectangular bar : Influence of the length on the reflection



$$1 = R^2 + T^2$$



shallow water

deep water

# Application to nearshore structures:

Expression of the velocity potential : presence of evanescent modes

$$\Phi(x, y, z, t) = \phi(x, z) e^{i\omega t}$$

$$= \left[ \left[ A^- e^{-ik_x x} + A^+ e^{+k_x x} \right] \chi(z) + \underbrace{\sum_{n=1}^{\infty} \left[ B_n^- e^{-k_{xn} x} + B_n^+ e^{+k_{xn} x} \right] \psi_n(z)}_{\text{evanescent}} \right] e^{i\omega t}$$

with

$$\chi(z) = \cosh k(h + z) = \psi_0(z) = \cos k_0(h + z)$$

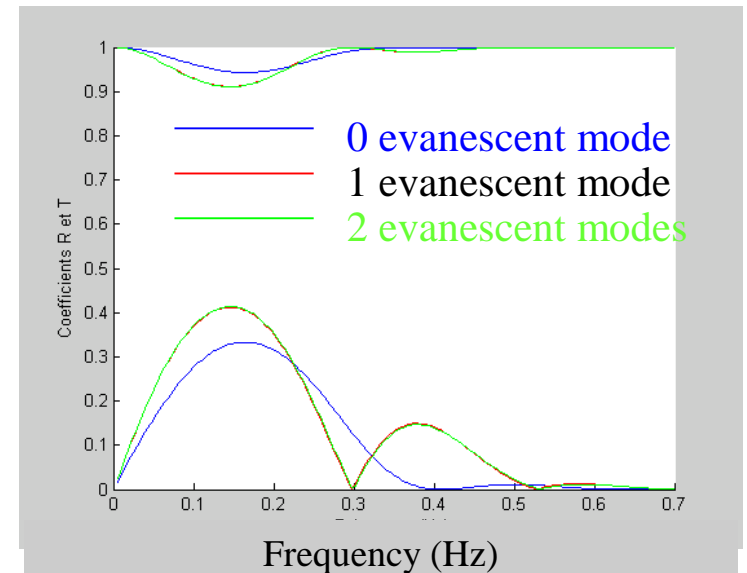
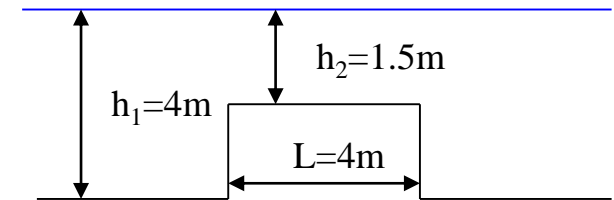
$$\psi_n(z) = \cos k_n(h + z)$$

Integral matching method, at a step:

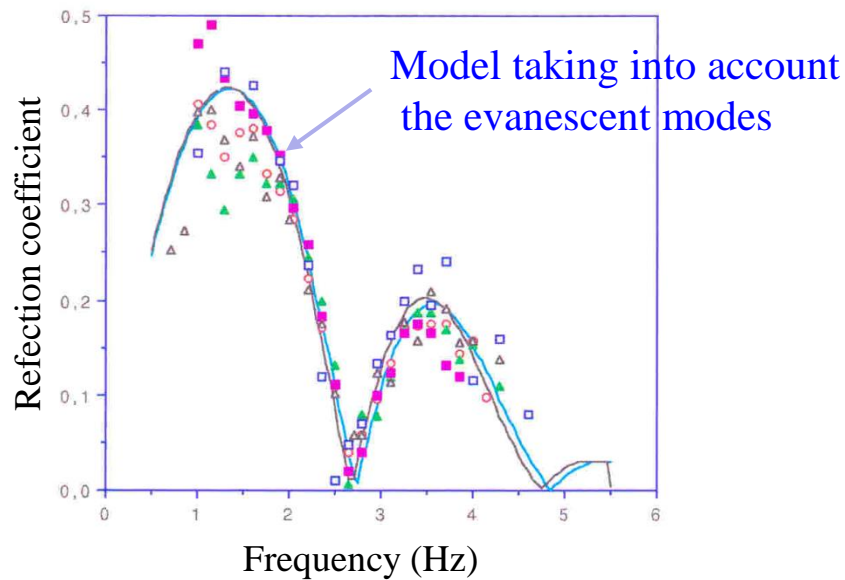
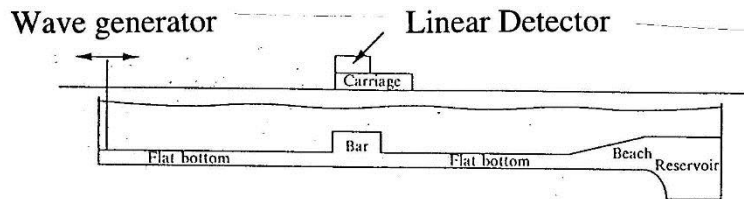
Pressure: 
$$\int_{-h_2}^0 \phi_1 \cdot \psi_{2,n} dz = \int_{-h_2}^0 \phi_2 \cdot \psi_{2,n} dz$$

Velocity: 
$$\int_{-h_1}^0 \frac{\partial \phi_1}{\partial x} \cdot \psi_{1,n} dz = \int_{-h_2}^0 \frac{\partial \phi_2}{\partial x} \cdot \psi_{1,n} dz$$

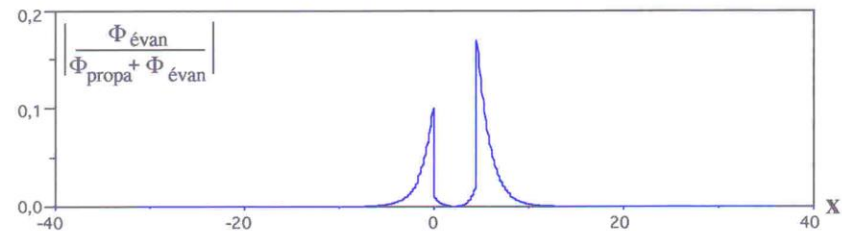
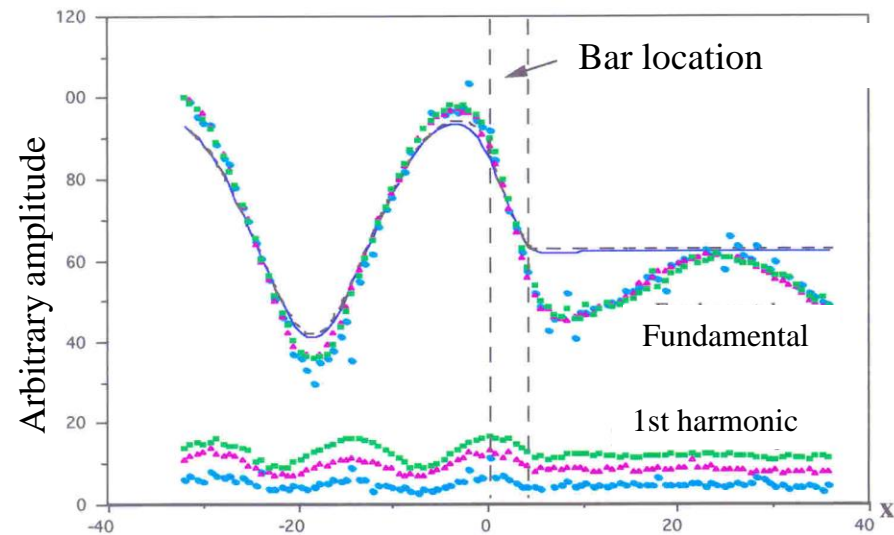
for  $n = 0, \dots, P$



# Weight of evanescent modes

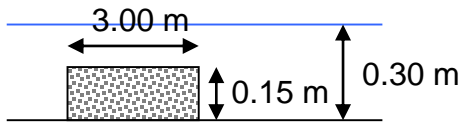


$f = 1 \text{ Hz}$

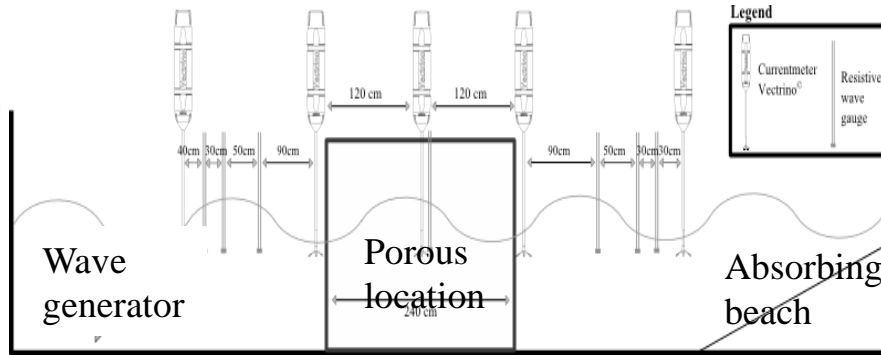


➔ Significant weight at the vicinity of the structure  
Evanescent modes involved in the interference process

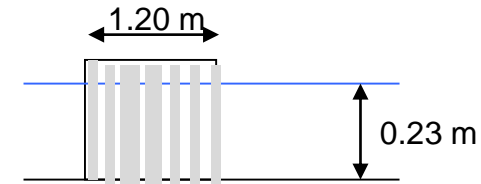
# Application to nearshore structures: Porous structures: reflection and dissipation



Porosity  $\gamma=0.3$



Experimental set-up, SeaTech wave tank, UTLN



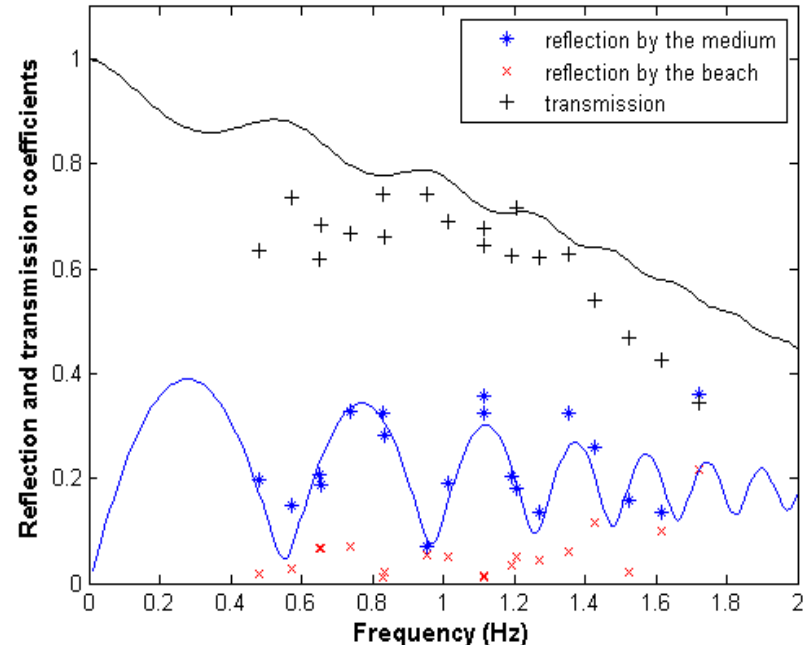
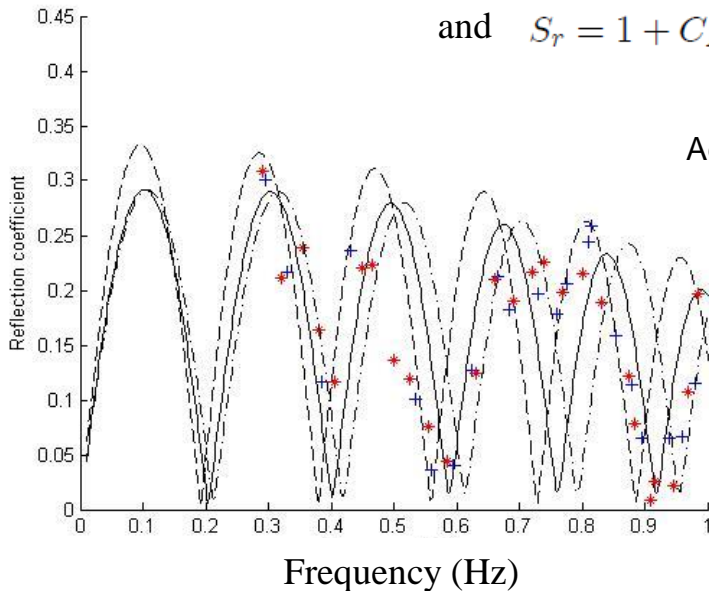
$D=0.050$  m  
Porosity  $\gamma=0.7$

Dispersion relation  $Z\omega^2 = igk \tanh(kh)$

with  $Z = f_R + iS$

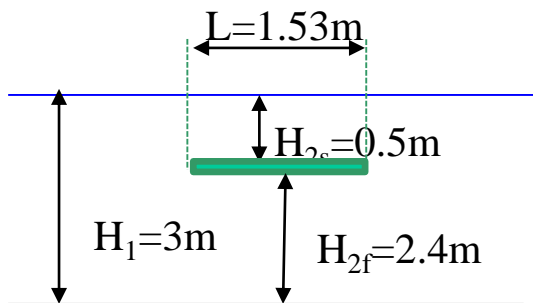
and  $S_r = 1 + C_M \frac{(1-\gamma)}{\gamma}$

Added mass

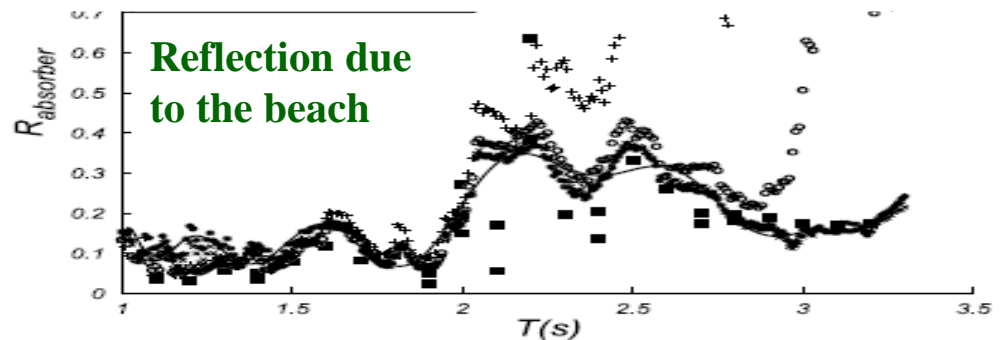
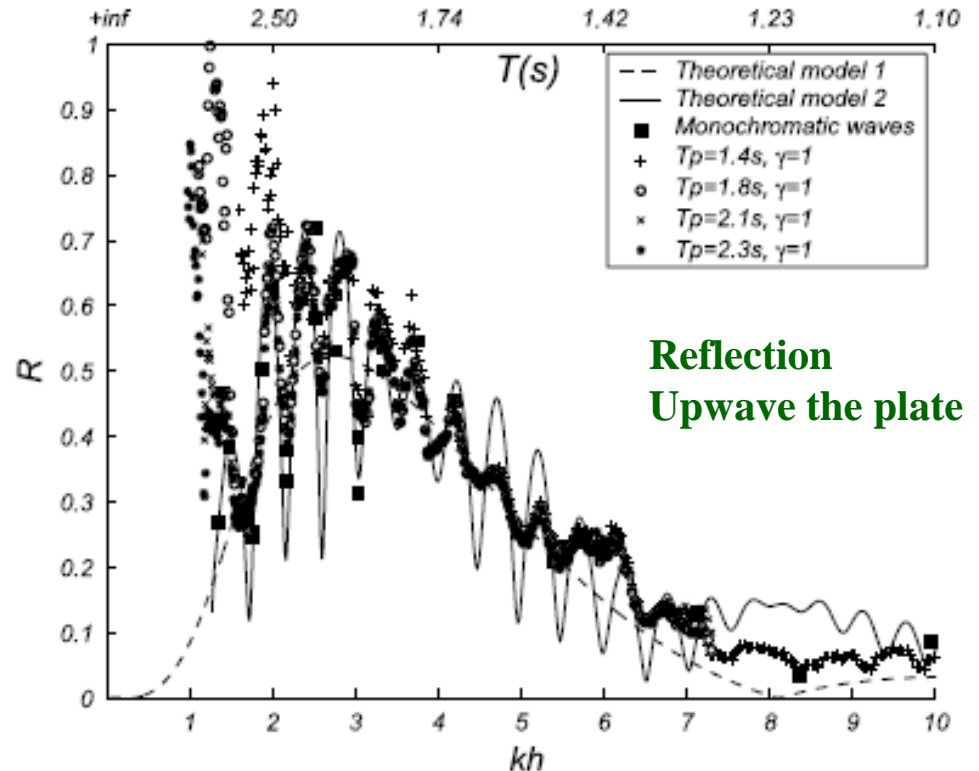


# Application to nearshore structures:

## Submerged plate : Interference process and Influence of the beach on the reflection



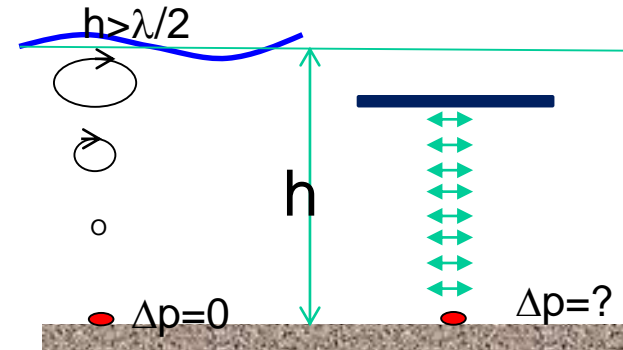
➔ Interference process above the plate  
Oscillating behaviour due to reflection from the beach (absorber)





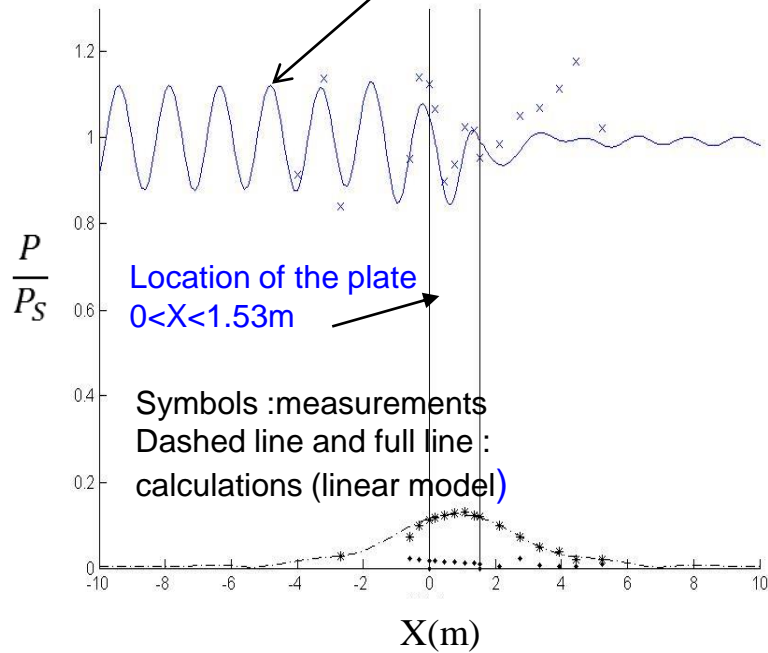
# Application to nearshore structures: Submerged plate : bottom induced pressure

Free surface deformation and bottom pressure amplitudes,  
T=1.4s, a=54mm

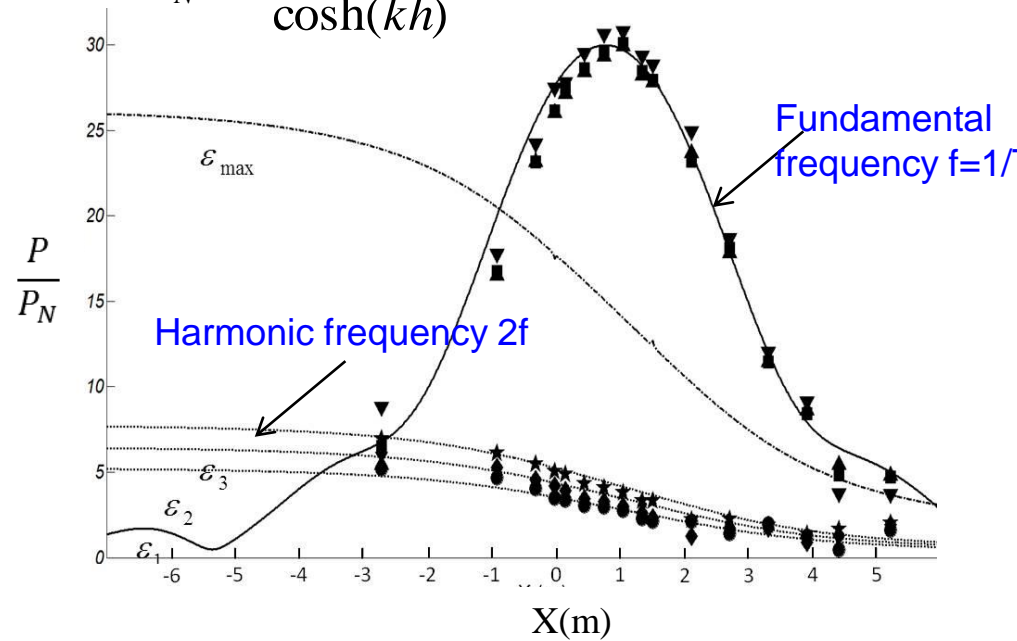


Partially standing wave

$$P_S = \rho g a_i$$



$$P_N = \frac{\rho g a_i}{\cosh(kh)}$$



- ➡ - 1st order bottom pressure amplitude (frequency  $f$ ) 30 times the bottom pressure due to the incoming wave
- ➡ - « Longuet-Higgins effect » at 2<sup>nd</sup> order (frequency  $2f$ )

# Application to wave energy device:

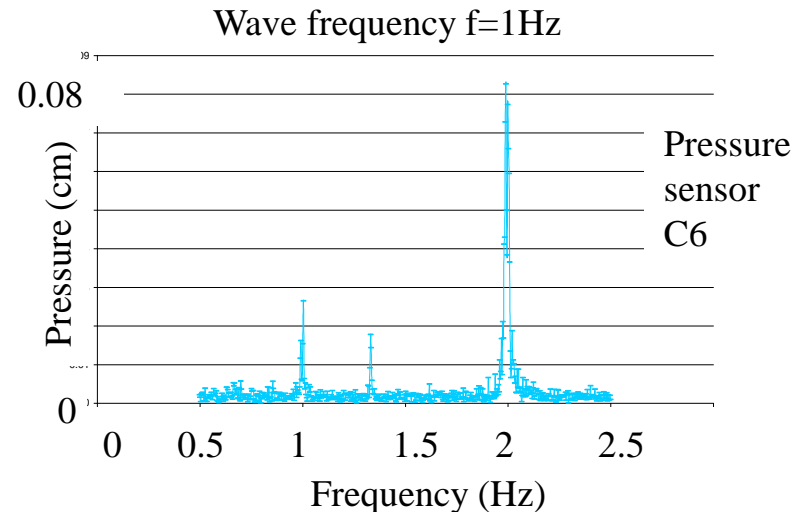
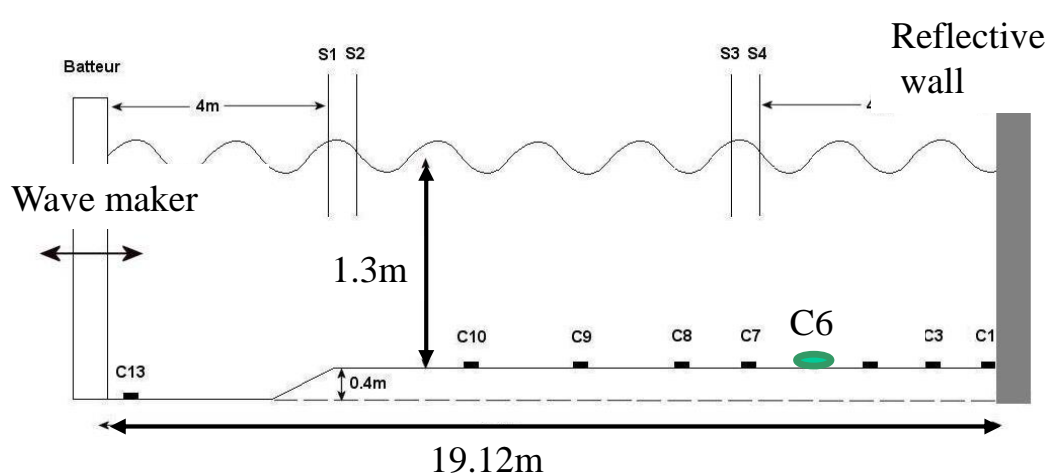
bottom induced pressure at twice the wave frequency for « deep water » conditions

Non linear free surface condition at  $z=\eta$ : 
$$g \frac{\partial \Phi}{\partial z} + \frac{\partial \vec{v}^2}{\partial t} + \frac{\partial^2 \Phi}{\partial t^2} + \frac{1}{2} \vec{v} \cdot \vec{\nabla} (\vec{v}^2) = 0$$

Perturbation method (for 2<sup>nd</sup> order Stokes waves): 
$$\Phi = \varepsilon \Phi_1 + \varepsilon^2 \Phi_2 + O(\varepsilon^2)$$

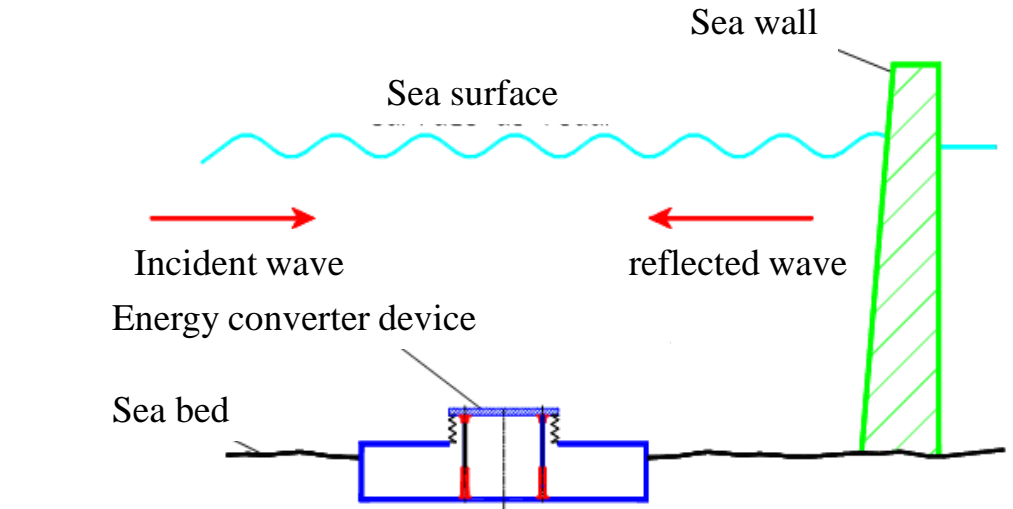
$$\begin{cases} \Phi_1 = \frac{a^- \omega}{k} e^{kz} \cos[\omega t - kx] + \frac{a^+ \omega}{k} e^{kz} \cos[\omega t + kx] \\ \Phi_2 = -a^- a^+ \sin[2\omega t] \end{cases}$$

$$p = -\rho \frac{\partial \Phi}{\partial t} - \frac{1}{2} \rho v^2 \quad \Rightarrow \quad p = \rho \omega a^- a^+ \cos(2\omega t) \quad \text{for } z=-h \text{ (in fact for any } z)$$

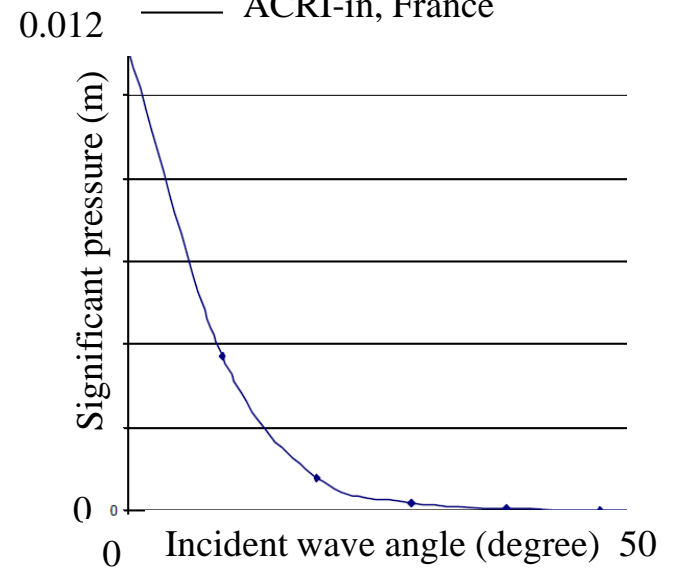
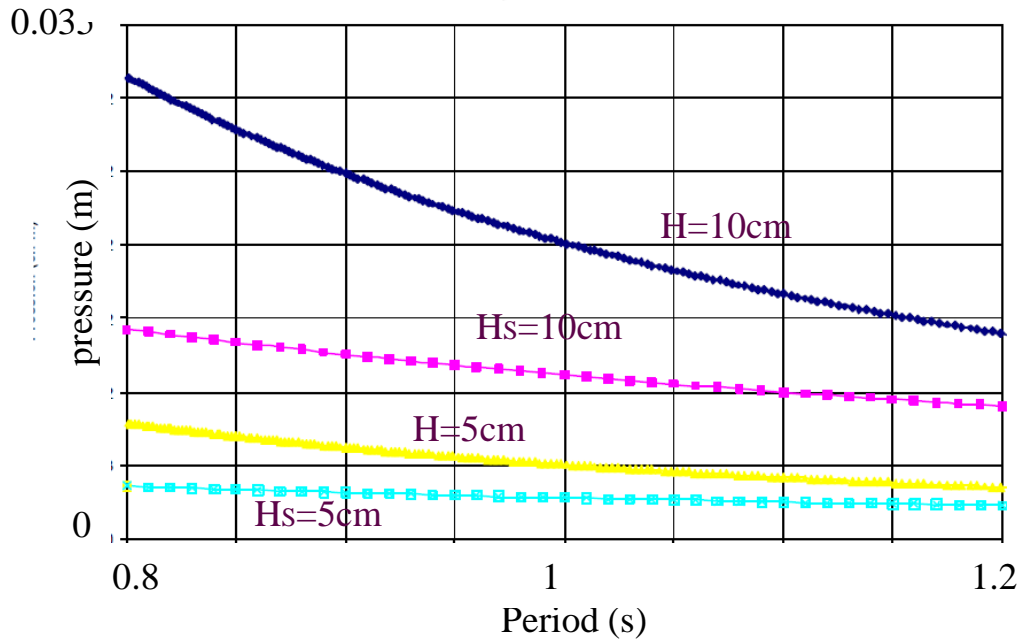


# Application to wave energy device:

## Real sea states



Experimental set-up,  
ACRI-in, France



➡ « Longuet-Higgins » effect significantly decreases for wave frequency and direction spreading

## Refraction – diffraction

Slow changes of wave properties :

Plane waves, ray theory

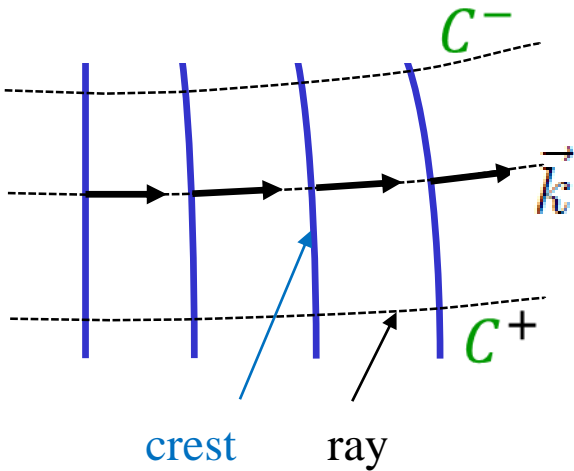
Rapid change of amplitude and/or direction :

No more plane waves, diffraction parameter

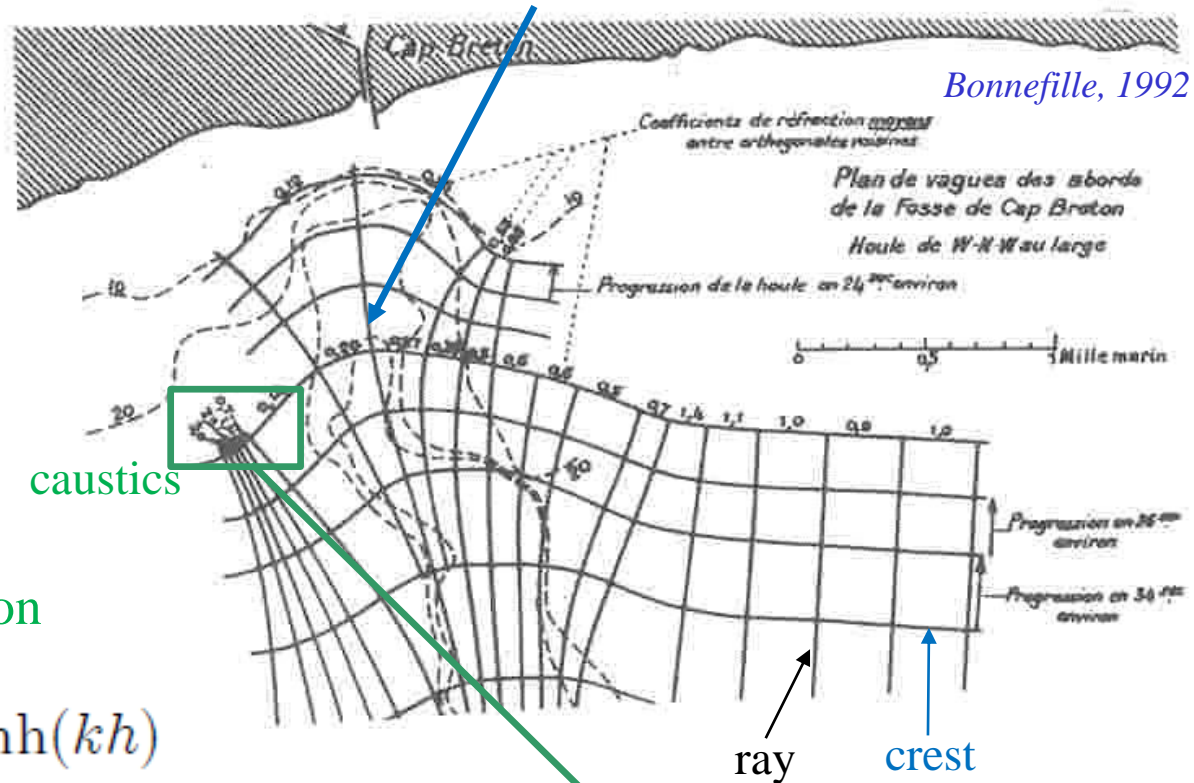


# Refraction – diffraction

## Wave celerity C

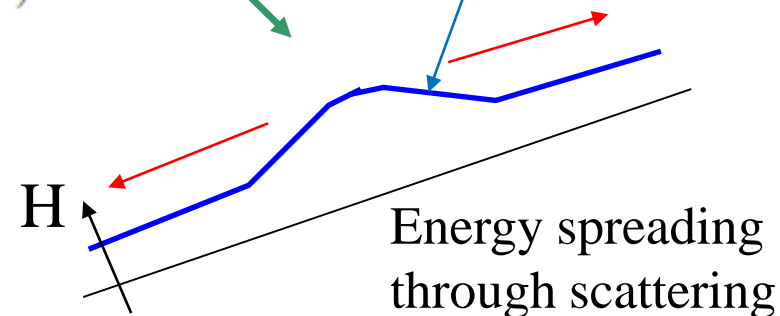


## Canyon of Capbreton, France



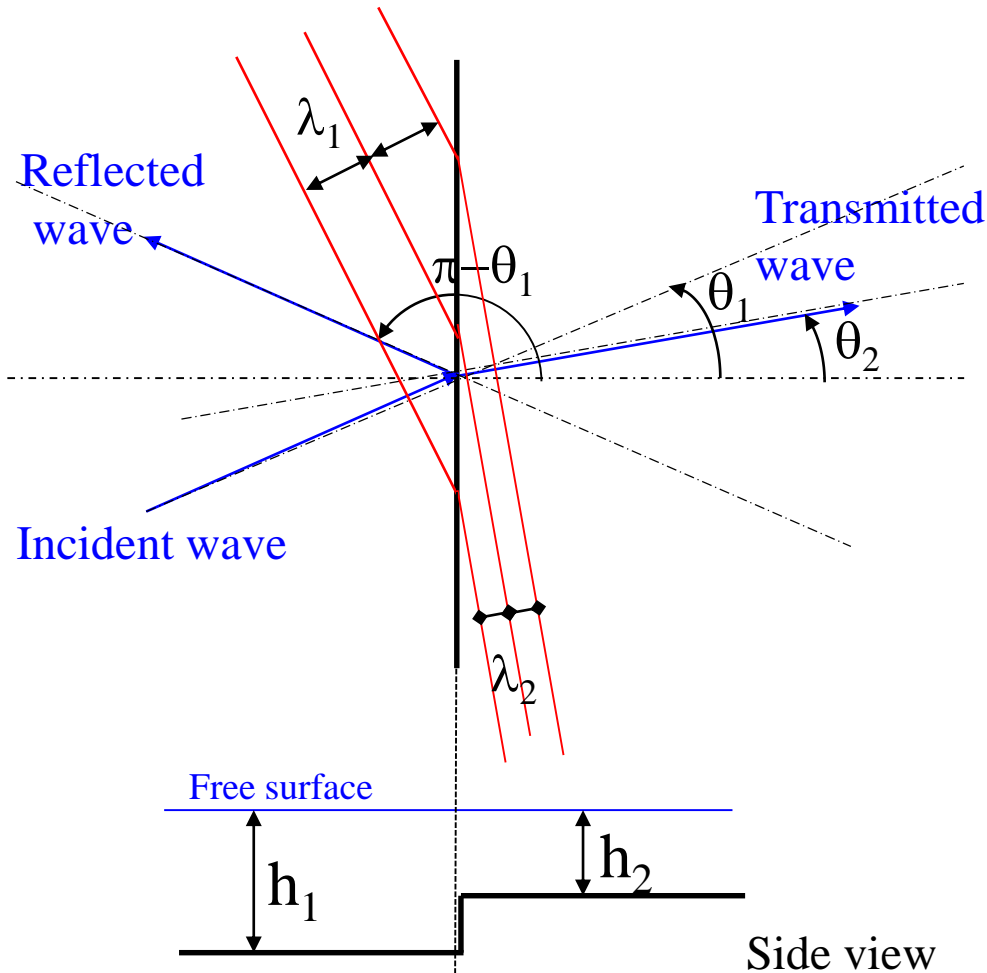
## Wave celerity $C = \omega/k$ depends on

- bathymetry :  $\omega^2 = gk \tanh(kh)$
- Currents :  $(\omega - Uk)^2 = \sigma^2 = gk \tanh(kh)$
- Porous media :  $Z\omega^2 = igk \tanh(kh)$



# Refraction – Descartes – Snell’s Law

1D step effect



Descartes-Snell’s Law

$$\frac{\sin \theta_1}{C_1} = \frac{\sin \theta_2}{C_2} \text{ with } C_1 = \frac{\omega}{k_1}, C_2 = \frac{\omega}{k_2}$$

Slowly varying depth  
(WKB approximation, mild slope)

$$\eta = a \exp(iS)$$

where  $S(\vec{x}, t, \vec{k}, \omega)$

in the plane wave approximation

$$\vec{k} = -\vec{\nabla}_h S \text{ and } \omega = \frac{\partial S}{\partial t}$$

$$\vec{k} (k_1 = k \cos \theta; k_2 = k \sin \theta)$$

and

$$S = \omega t - k_1 x - k_2 y$$

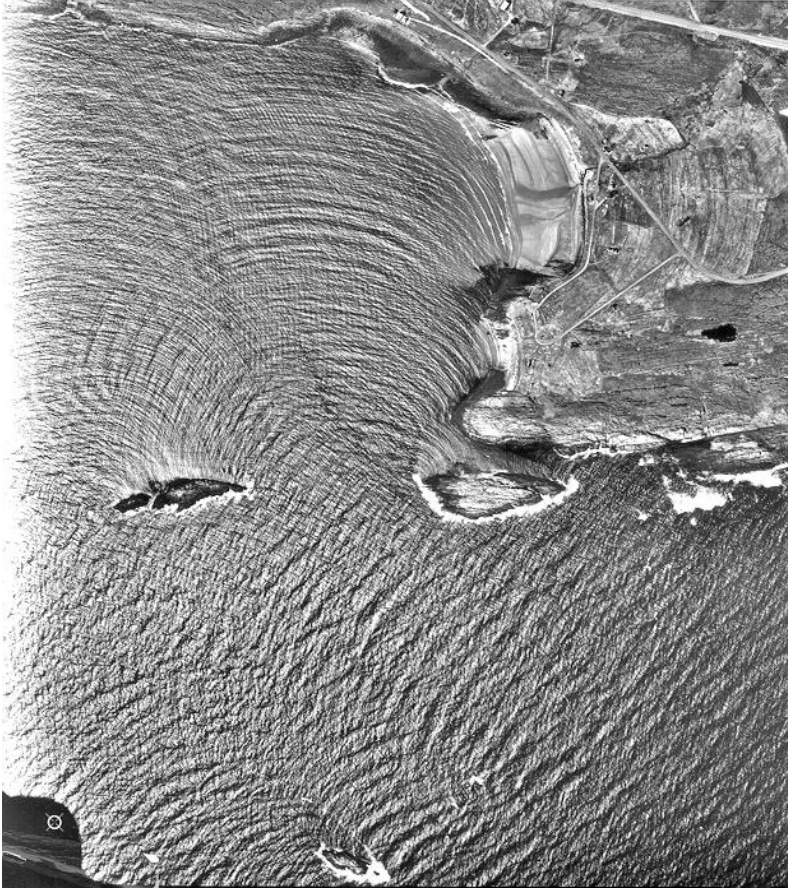
Eikonal equation

$$\left(\frac{\partial S}{\partial x}\right)^2 + \left(\frac{\partial S}{\partial y}\right)^2 = \left(\vec{\nabla}_h S\right)^2 = k^2$$

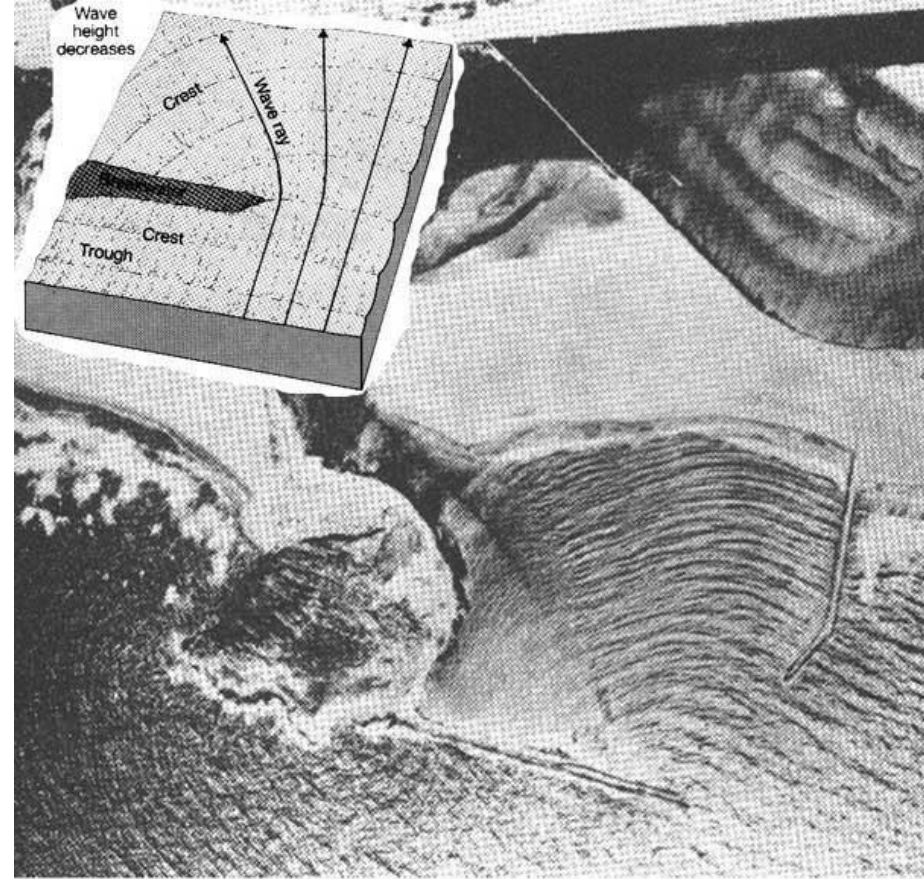
since  $\overrightarrow{\text{curl}}(\vec{k}) = \vec{0}$ :

$$\frac{\partial(k \sin \theta)}{\partial x} - \frac{\partial(k \cos \theta)}{\partial y} = 0$$

# Diffraction



Diffraction due to rapid bathymetric changes

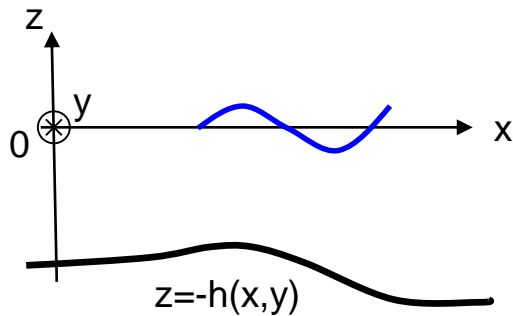


Diffraction occurs behind the breakwater.

Diffraction from abrupt structures (dikes, jetties)

# Refraction-diffraction due to « rapid » bathymetric changes

## Berkhoff equation



Velocity potential :  $\Phi(x, y, z, t) = \phi(x, y, z)e^{i\omega t}$

$$\phi(x, y, z) = \frac{ig \cosh[k(z+h)]}{\omega \cosh(kh)} \eta(x, y)$$

Hypothesis (mild slope) :  $\mu = \frac{\nabla h}{kh} \ll 1$

Conservation equation and boundary conditions :

$$\frac{\partial^2 \phi}{\partial z^2} = -\nabla_h^2 \phi \text{ for } 0 \geq z \geq -h(x, y)$$

$$\frac{\partial \phi}{\partial z} - \frac{\omega}{g} \phi = 0 \text{ for } z = 0$$

$$\frac{\partial \Phi}{\partial z} = \vec{\nabla}_h h \vec{\nabla}_h \phi \text{ for } z = -h(x, y)$$

Weak formulation of the Laplace's equation :

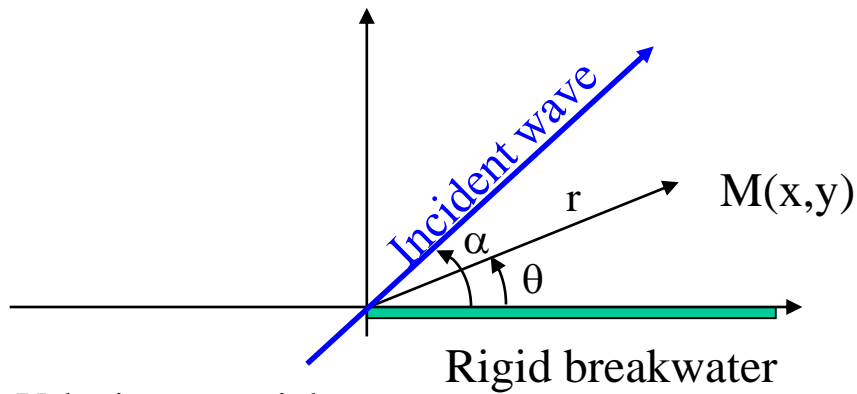
$$\int_{-h}^0 (k^2 f \phi + \nabla_h^2 \phi f) dz = [(\nabla_h h \nabla_h \phi) f]_{z=-h} \text{ with the Green function } f = \cosh[k(z+h)]$$

Mei, 1983  $\longrightarrow$   $\vec{\nabla}_h \cdot (CC_g \vec{\nabla}_h \eta) + k^2 CC_g \eta = 0$  **Berkhoff equation**



# Diffraction due to vertical walls : **Diffraction behind a breakwater**

Wave of incidence  $\alpha$  :



Velocity potential :

$$\Phi(x, y, z, t) = \frac{\omega}{k} \frac{\cosh[k(z+h)]}{\sinh(kh)} F(x, y) e^{i\omega t}$$

F verifies the Helmholtz equation :  $\nabla_h^2 F + k^2 F = 0$

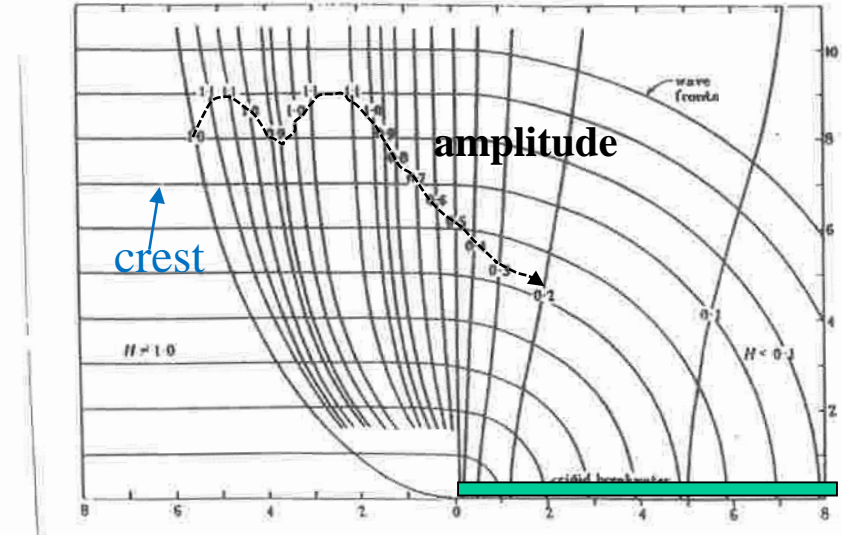
Solution (polar coordinates) :

$$F(r, \theta) = I \left( -\sqrt{\frac{4kr}{\pi}} \sin \frac{\alpha - \theta}{2} \right) e^{ikr \cos(\alpha - \theta)}$$

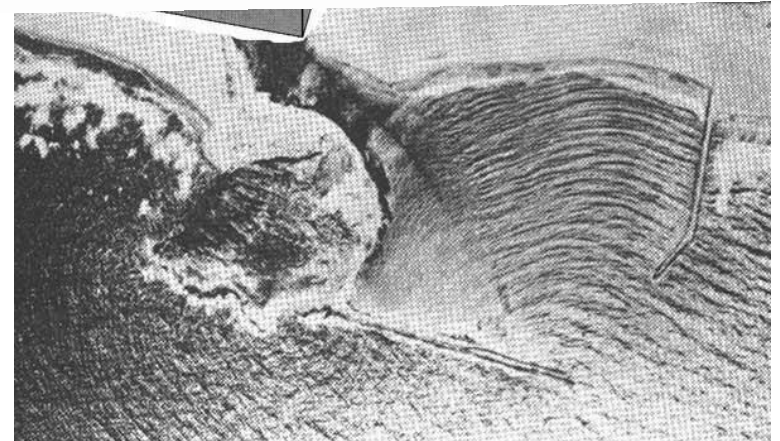
$$+ I \left( -\sqrt{\frac{4kr}{\pi}} \sin \frac{\alpha + \theta}{2} \right) e^{-ikr \cos(\alpha + \theta)}$$

with  $I(u) = \frac{1}{2}(1+i) \int_{-\infty}^u e^{-i\frac{\pi u'}{2}} du'$

Normal incidence ( $\alpha = \pi/2$ ) :



Rigid breakwater



Wave diffraction behind the breakwater

# Propagation in the presence of vertical walls : Expression of the velocity potentials

$$\begin{aligned}\Phi(x, y, z, t) &= \cosh[k(z + h)]\phi(x, y)e^{i\omega t} \\ &= \cosh[k(z + h)] \sum_{n=0}^{\infty} \left[ A_n^- e^{-ik_{xn}x} + A_n^+ e^{+ik_{xn}(x-L)} \right] \psi_n(y) e^{i\omega t}\end{aligned}$$

with  $\psi_n(y) = \cos k_{yn}(y - d_m)$

where  $k_{yn} = \frac{n\pi}{d_M - d_m}$

Since  $\nabla^2 \Phi = 0$ ,  $k_{xn} = (k^2 - k_{yn}^2)^{\frac{1}{2}}$

$$n_{prop} = 1 + \text{Int} \left[ \frac{kd}{\pi} \right]$$

$n < n_{prop}$  Propagating modes of direction  $\theta_n = \pm \text{atan} \frac{|k_{yn}|}{|k_{xn}|}$

$n > n_{prop}$  Evanescent modes along the x-axis

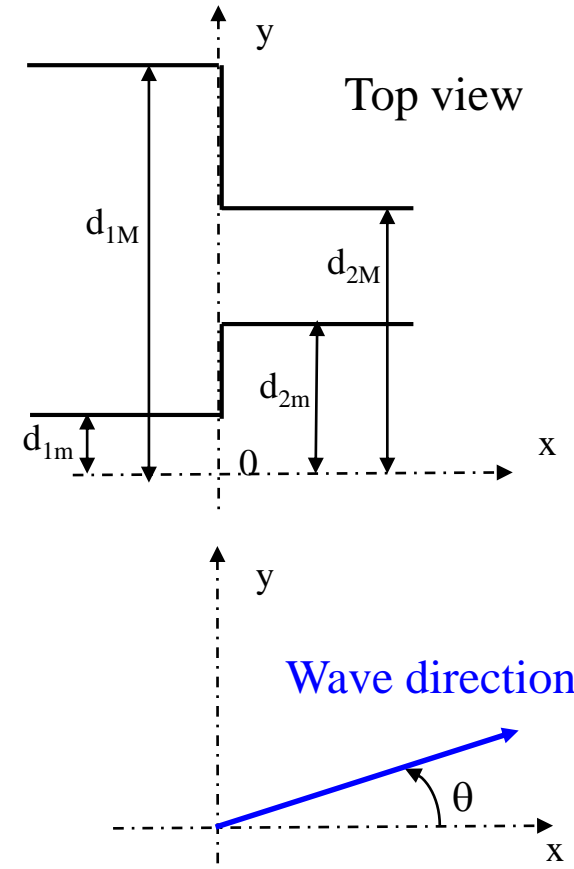
Integral matching method, at a width discontinuity:

Pressure: 
$$\int_{d_{m2}}^{d_{M2}} \phi_1 \cdot \psi_{2,n} dy = \int_{d_{m2}}^{d_{M2}} \phi_2 \cdot \psi_{2,n} dy$$

Velocity: 
$$\int_{d_{m1}}^{d_{M1}} \frac{\partial \phi_1}{\partial x} \cdot \psi_{1,n} dy = \int_{d_{m2}}^{d_{M2}} \frac{\partial \phi_2}{\partial x} \cdot \psi_{1,n} dy$$

for  $n = 0, \dots, P$

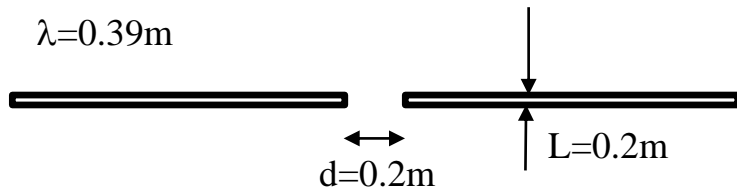
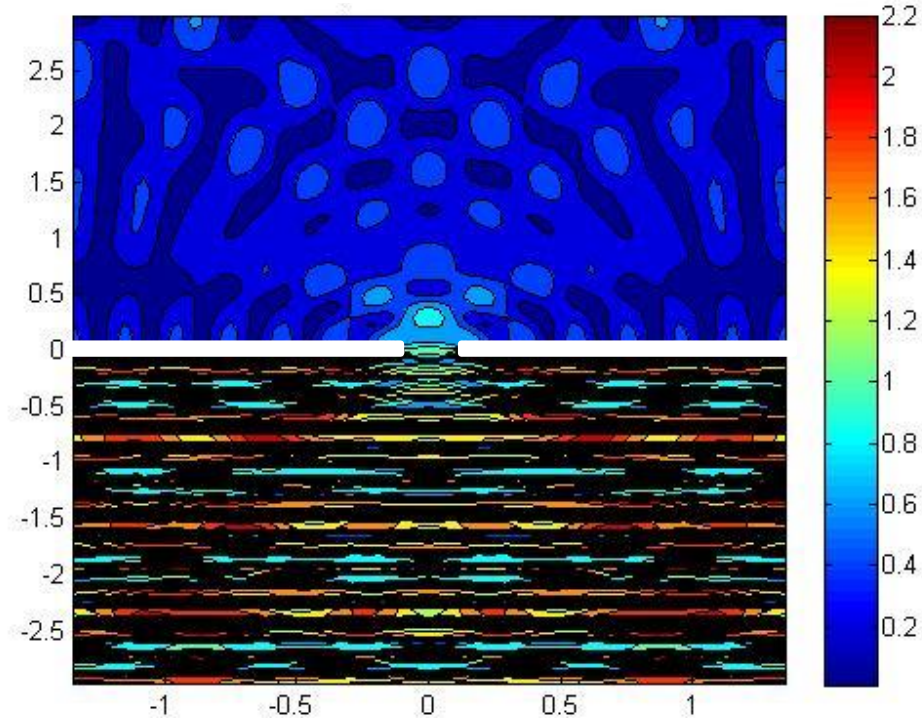
$h = \text{constant water depth}$



# Propagation in the presence of vertical walls : examples

## Sub-wavelength hole

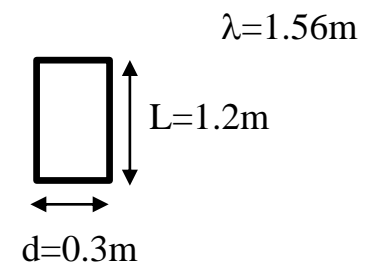
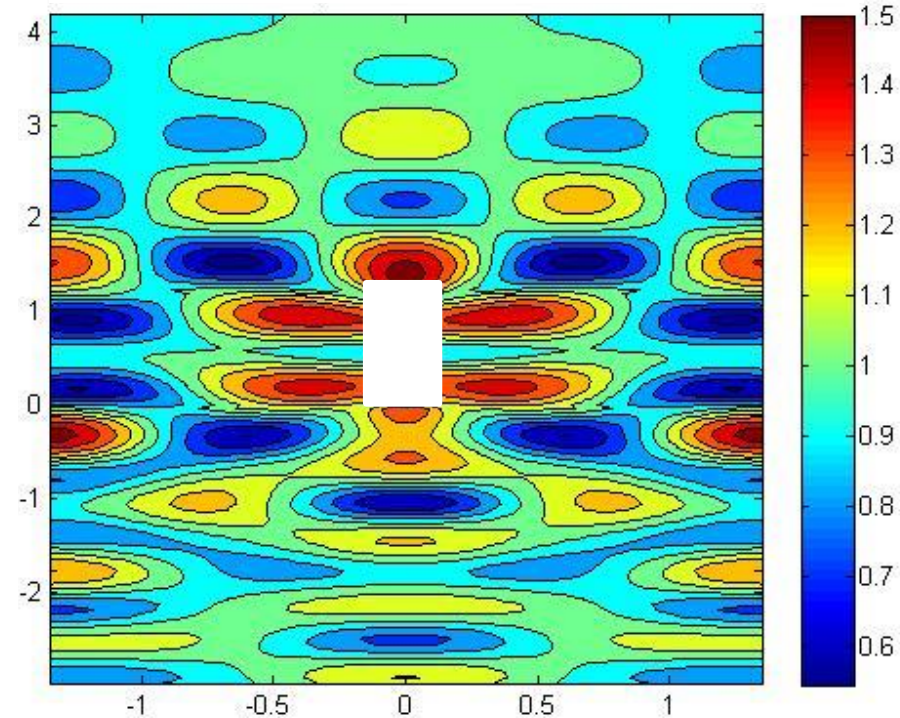
dimensionless amplitude with respect to the incoming wave



## Rectangular emerging structure

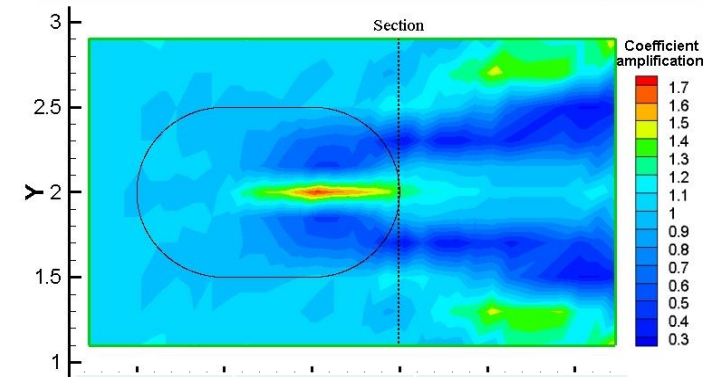
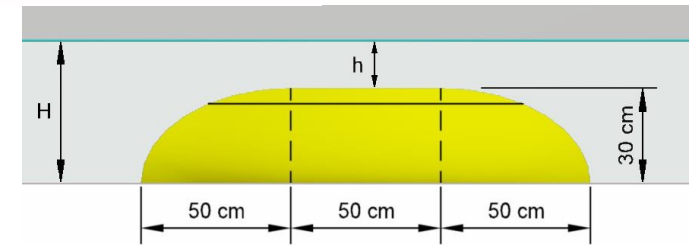
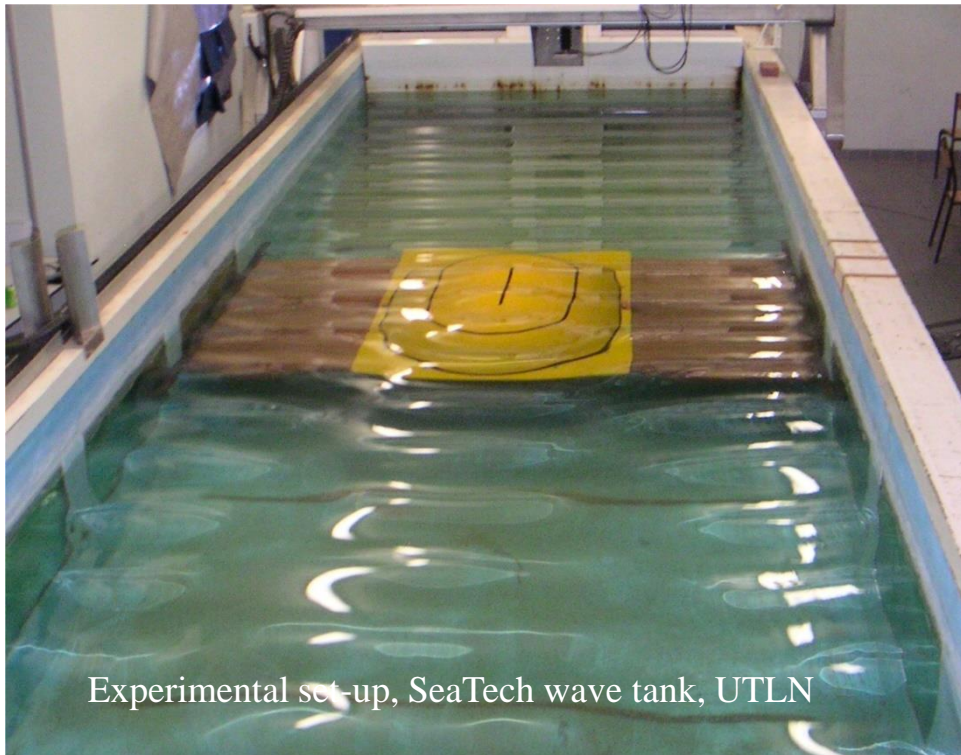
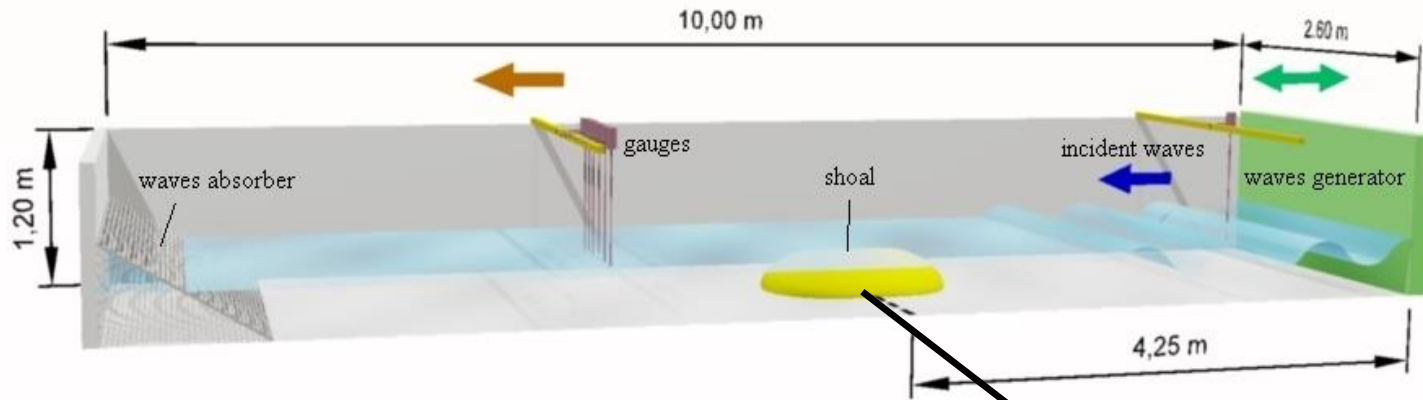
Top views

dimensionless amplitude with respect to the incoming wave



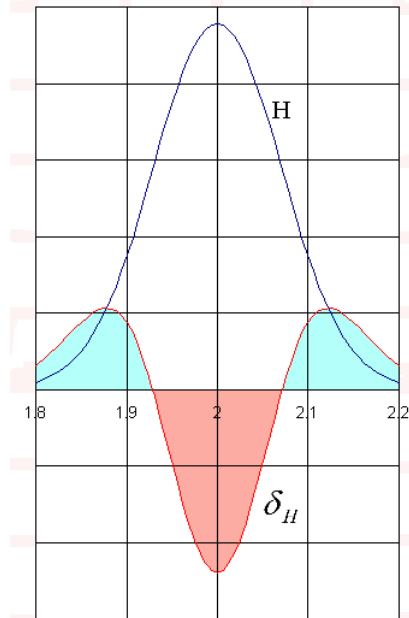
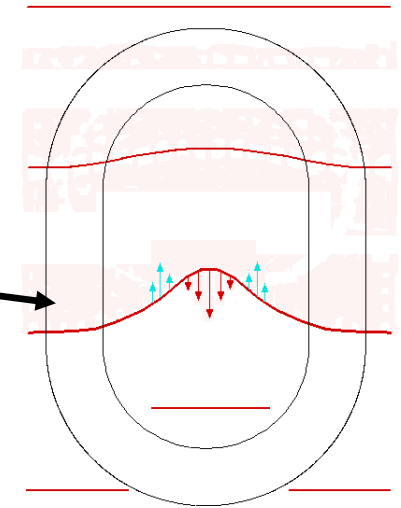
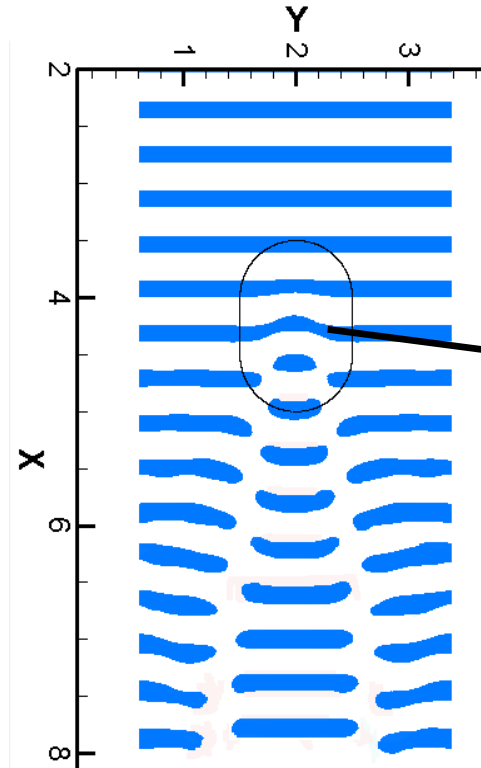
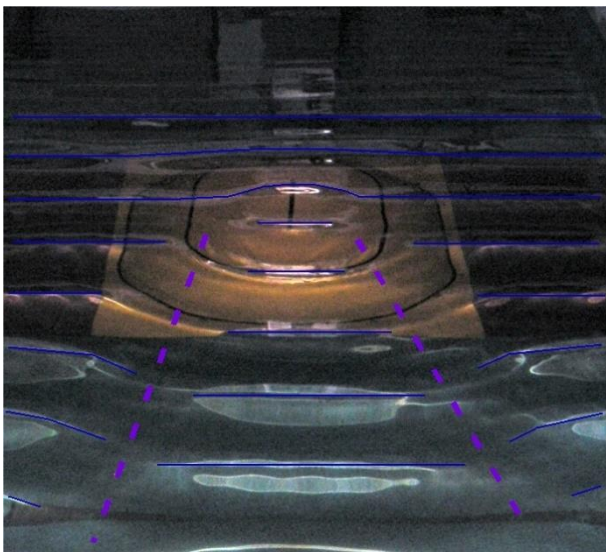
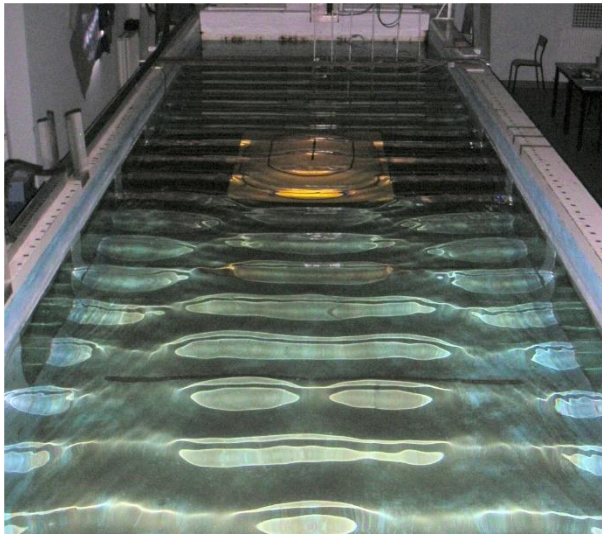
➡ either spreading or focusing effects

# Propagation in inhomogeneous media : wave focusing above a shoal



Period	Maximum of amplification	Distance from the end of the mound (m)
T=0.3	1	-
T=0.4	1.12	-0.20
T=0.5	1.31	-0.10
T=0.6	1.53	-0.10

# Wave focusing above a shoal : influence of the diffraction term $\delta_H$

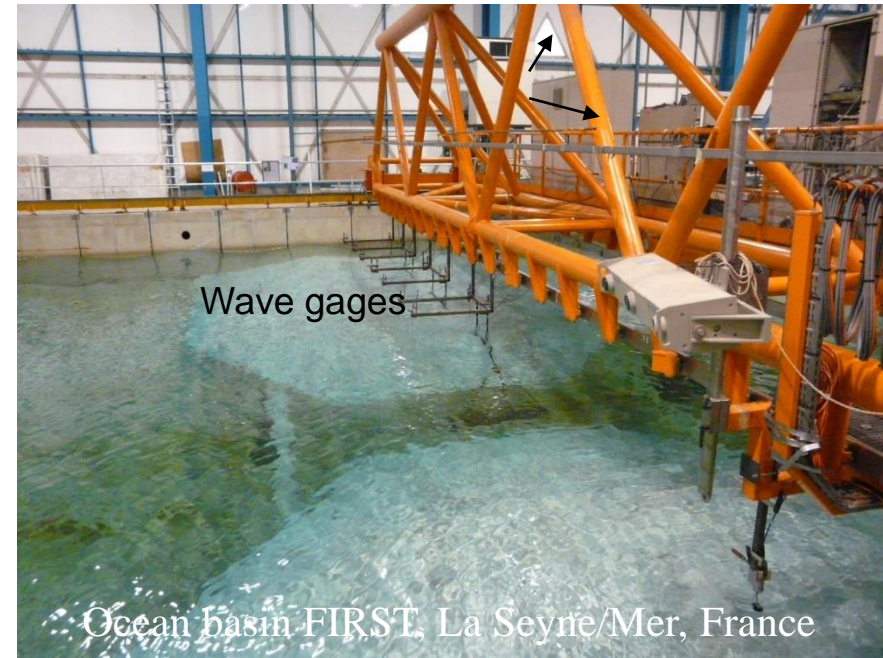
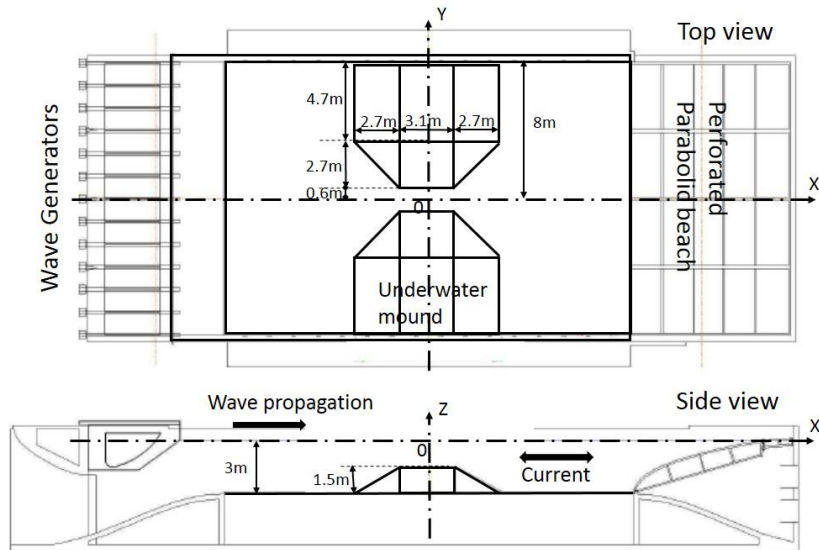


$$H = \hat{H} e^{iS}$$

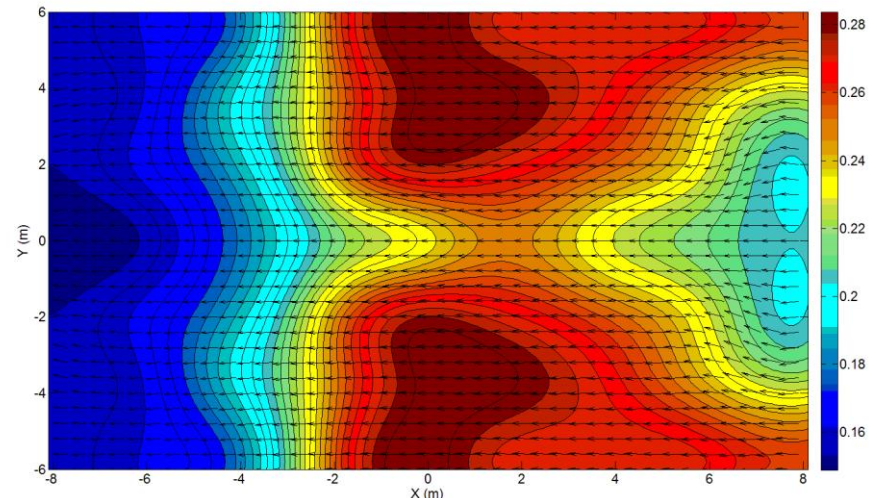
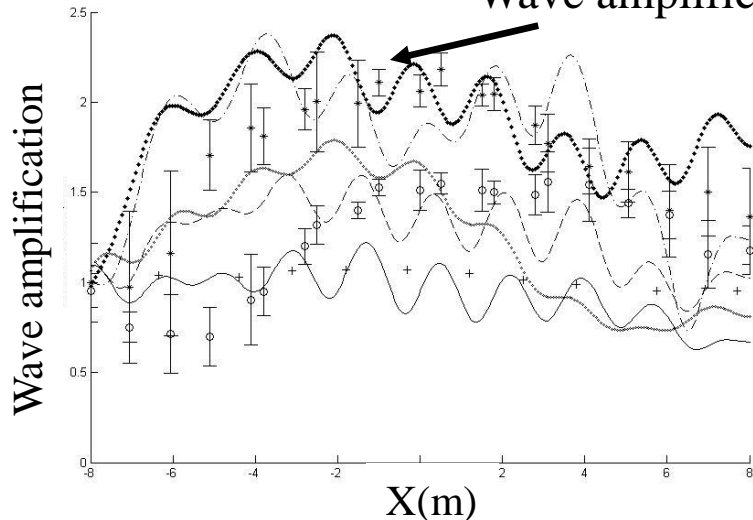
$$\delta_H = \frac{\vec{\nabla}_h \cdot (CC_g \hat{H}^2 \vec{\nabla}_h S)}{k^2 CC_g \hat{H}}$$

$$\begin{cases} k' = k\sqrt{1 + \delta_H} \\ C' = C/\sqrt{1 + \delta_H} \\ C'_g = C_g/\sqrt{1 + \delta_H} \end{cases}$$

# Propagation in inhomogeneous media : wave focusing in the presence of currents



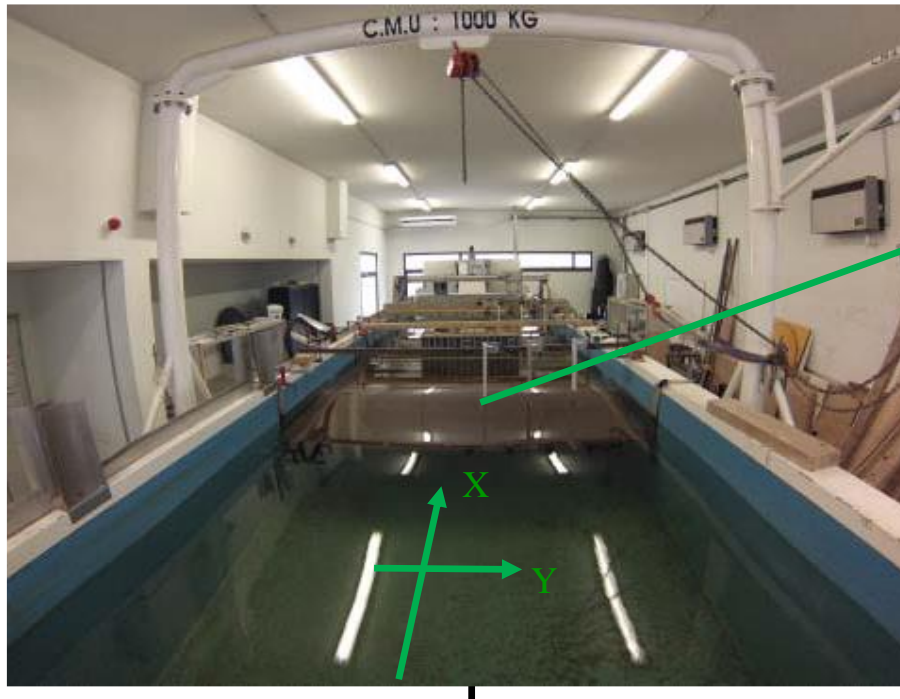
## Wave amplification



Mean current field

- ➔ current field forced by the underwater mounds
- ➔ Amplification up to two times the incident wave amplitude in deep water conditions for wave-opposing current conditions

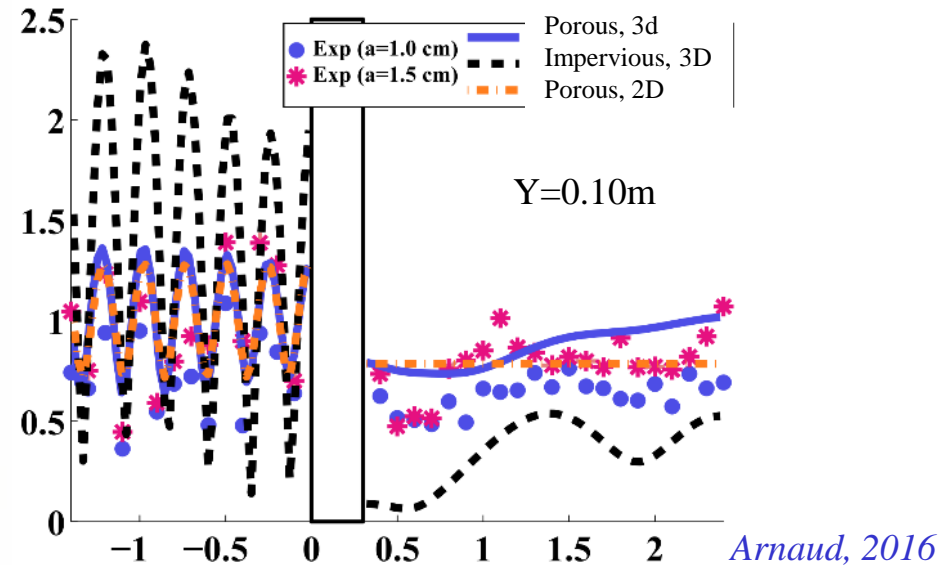
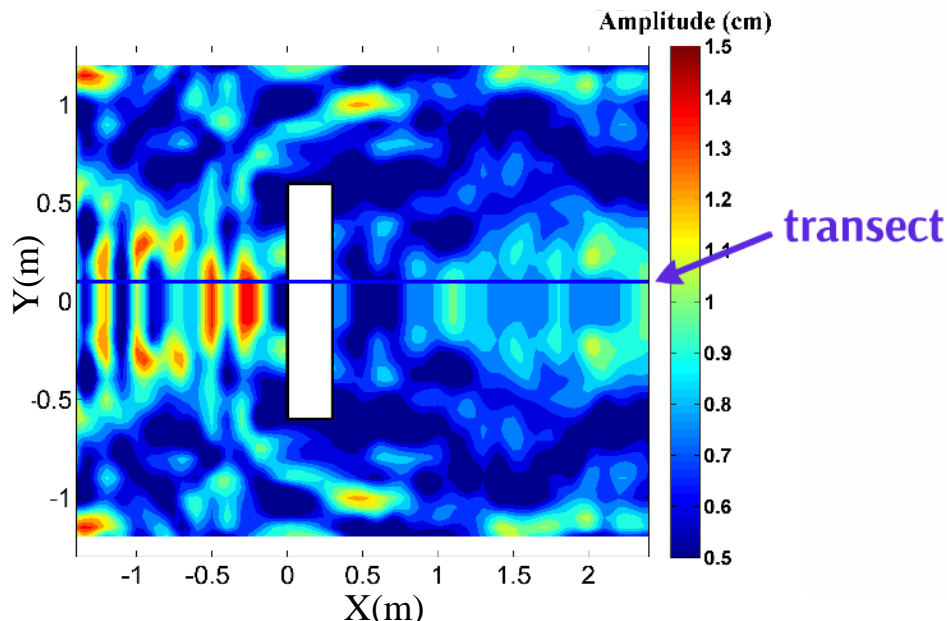
# Propagation in inhomogeneous media : wave scattering in the presence porous media



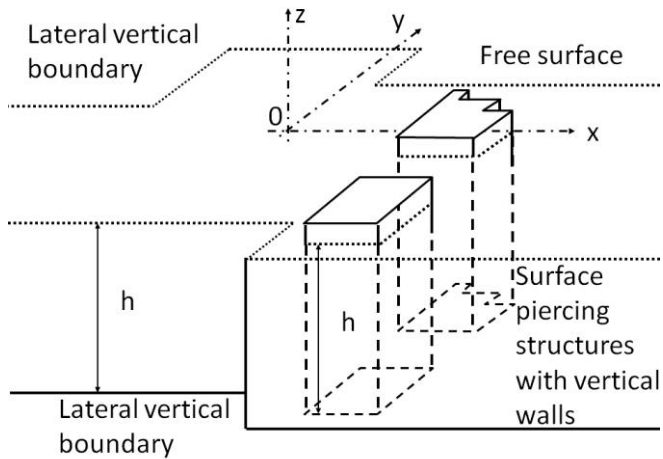
Porous rectangular structure  
made of vertical cylinders  
 $L \times l = 0.30\text{m} \times 1.20\text{m}$



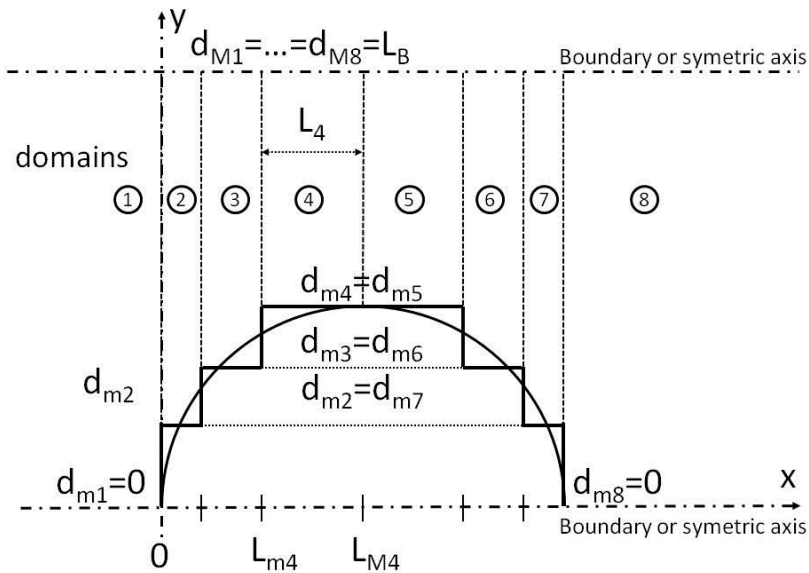
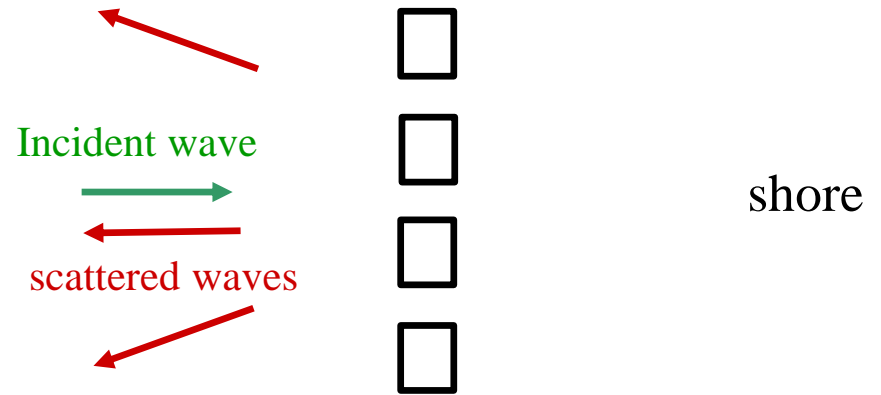
➔ wave scattering:  
Reflection upwave, refraction/diffraction  
propagation across the porous medium



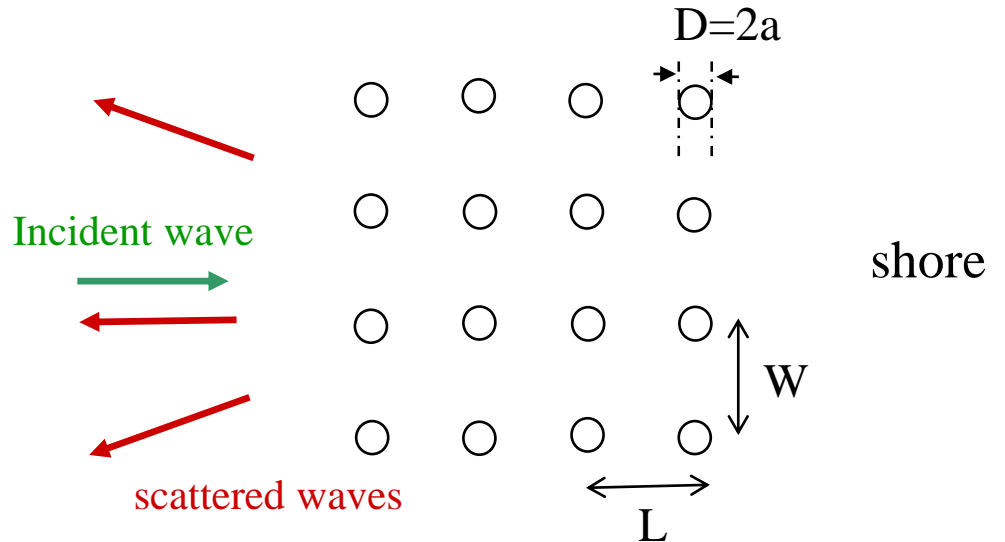
# Propagation in the presence of periodic structures : wave diffraction and Bragg resonance



➔ Periodic breakwaters of rectangular shape



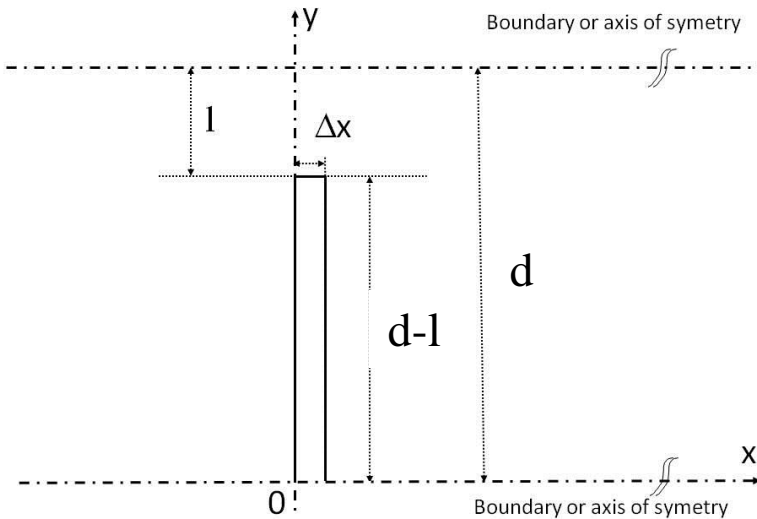
➔ Periodic cylinder arrays





# Periodic breakwaters of rectangular shape

## wave scattering



$k$ , wave wavenumber,  
Periodicity  $d$  of the breakwaters along  $y$ -axis

Number of propagating modes:

$$n_{prop} = 1 + \text{Int} \left[ \frac{kd}{\pi} \right]$$

### Reflection coefficient versus $kb$ :

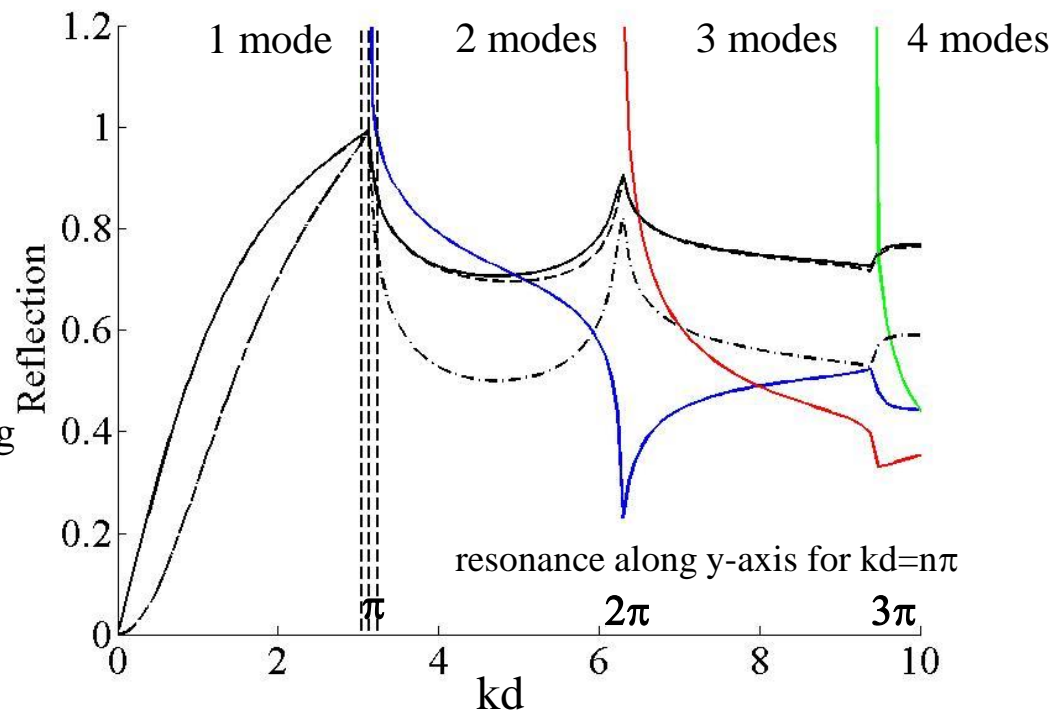
(-), 1st mode ( $n=0$ ); (- -) second mode ( $n=1$ );  
(- - -) third mode ( $n=2$ ); (- - - -) fourth mode ( $n=3$ );

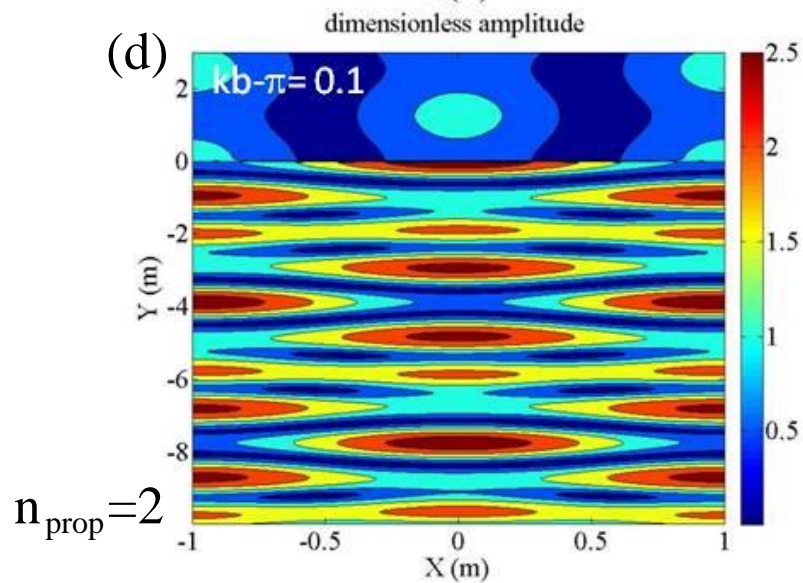
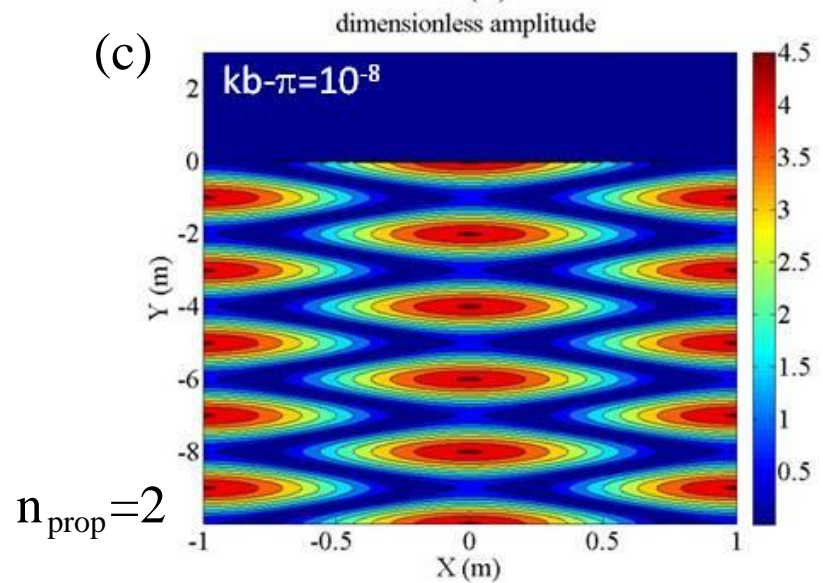
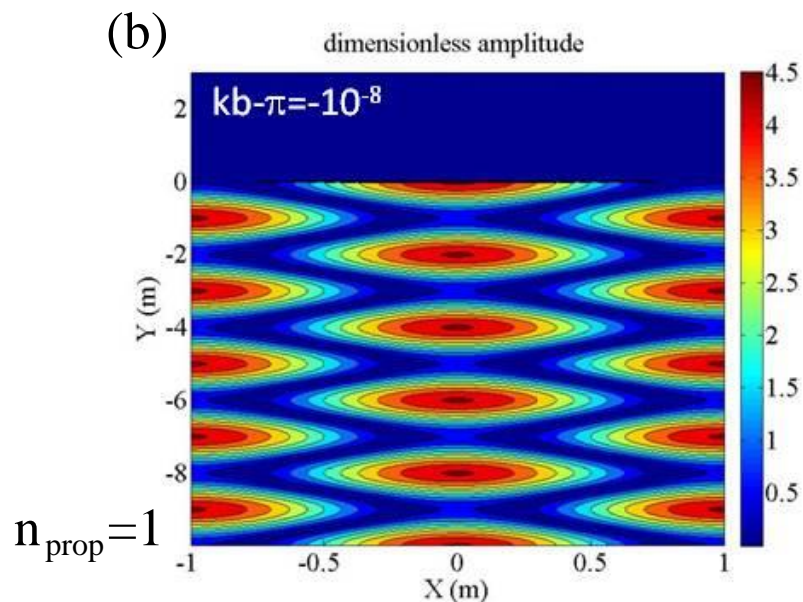
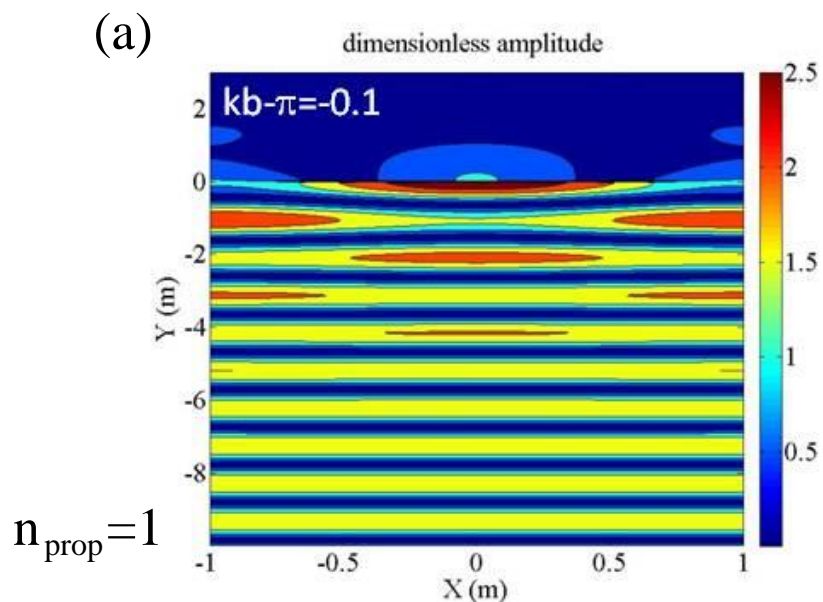
### Relative reflected energy flux versus $kb$ :

(-.-), 1st mode ( $n=0$ ), (- - -), total energy. Vertical dashed lines correspond to the locations of the frequency cases presented in the following

### Energy flux conservation:

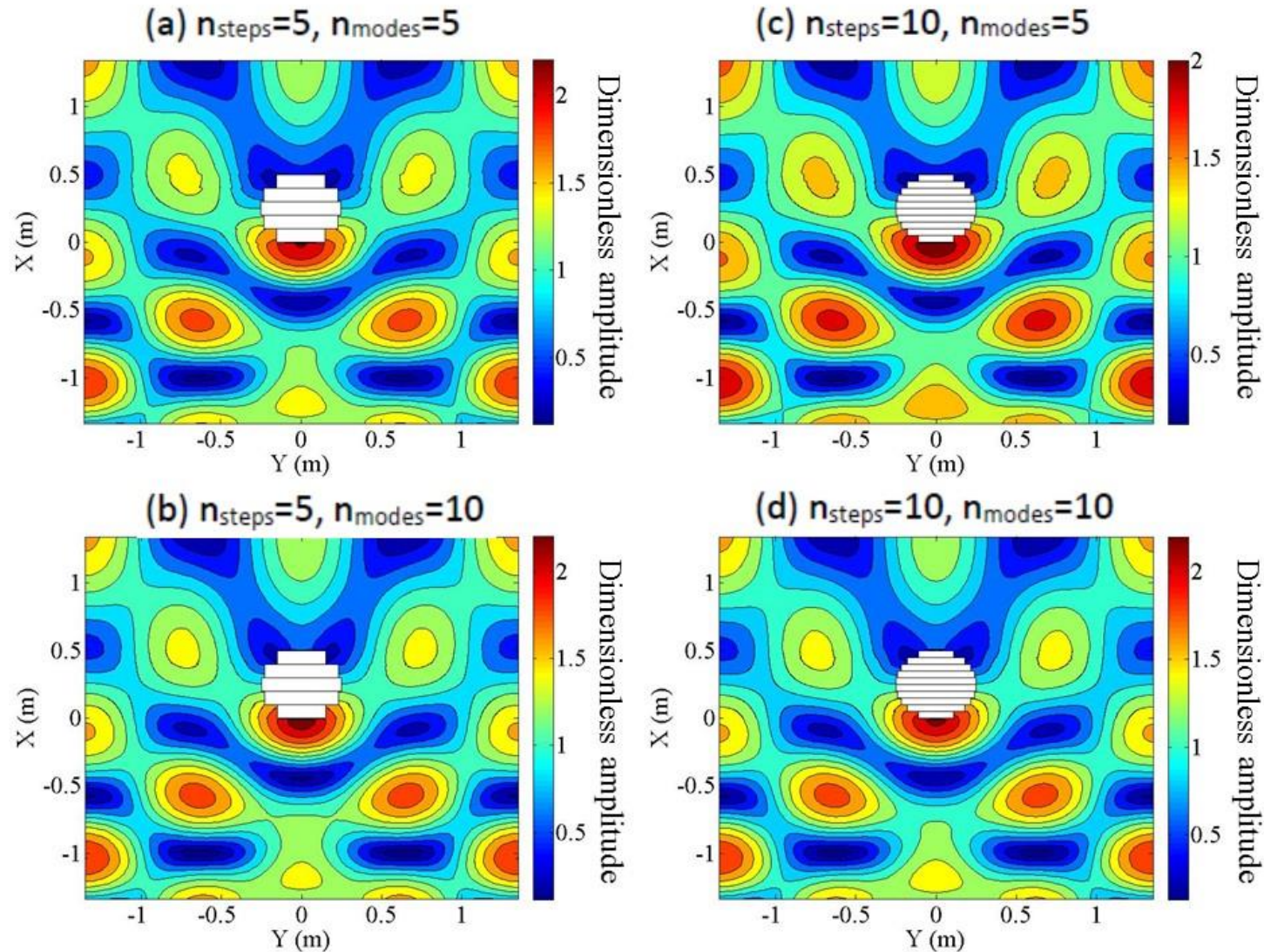
$$1 = R_0^2 + T_0^2 + \frac{1}{2} \sum_{n=1}^{n_{prop}-1} \frac{k_{xn}^{(0)}}{k} [R_n^2 + T_n^2]$$





Relative amplitude field relative to the incoming wave amplitude,  
 (a)  $kb - \pi = -0.1$ , (b)  $kb - \pi = -10^{-8}$ , (c)  $kb - \pi = 10^{-8}$ , (d)  $kb - \pi = 0.1$

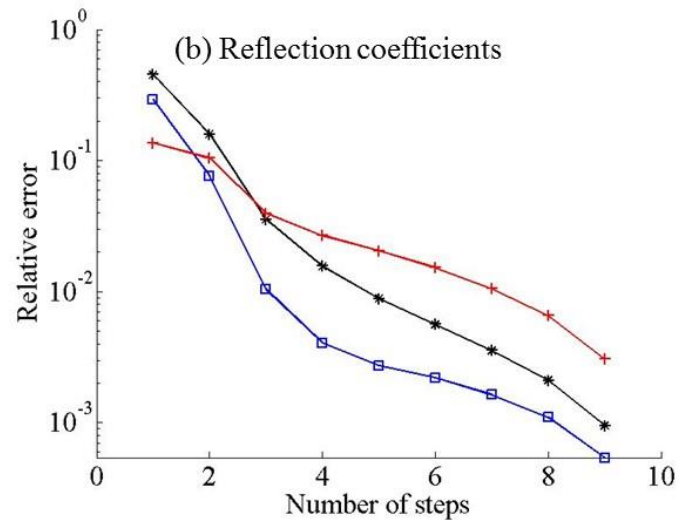
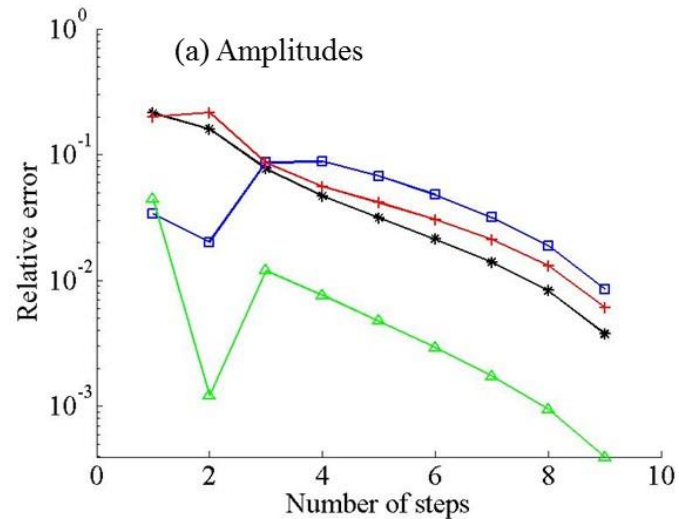
Propagation in the presence of periodic cylinders :  
Convergence with respect to the numbers of modes and steps



Relative surface wave amplitude field (a),  $n_{steps} = 5, n_{modes} = 5$ ; (b),  $n_{steps} = 5, n_{modes} = 10$ ; (c),  $n_{steps} = 10, n_{modes} = 5$ ; (d),  $n_{steps} = 10, n_{modes} = 10$

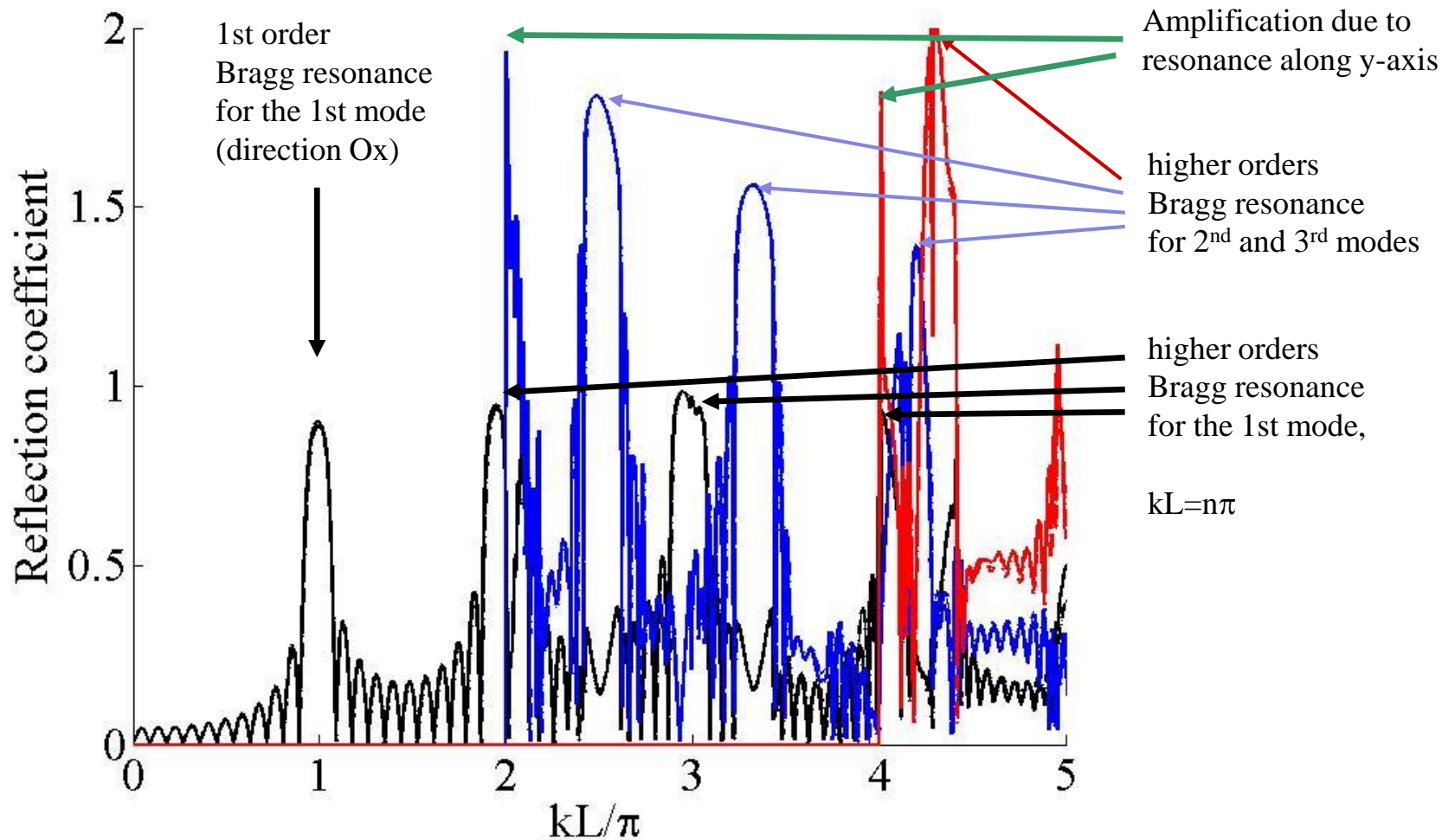
# Propagation in the presence of periodic cylinders :

## Convergence with respect to the number of steps ( for 10 modes)



Relative error (a) of the amplitudes, (b) of the reflection coefficients versus the number of steps  $n_{steps}$ . (a), (-\*),  $x = -0.0135$ ,  $y = 0$ ; (-□),  $x = 0.5085$ ,  $y = 0$ ; (-+),  $x = -0.0135$ ,  $y = 0.6$ ; (-△),  $x = 0.5085$ ,  $y = 0.6$ . (b) (-\*), first mode; (-□), second mode; (-+), third mode.

# Propagation in the presence of periodic cylinders : Bragg resonances



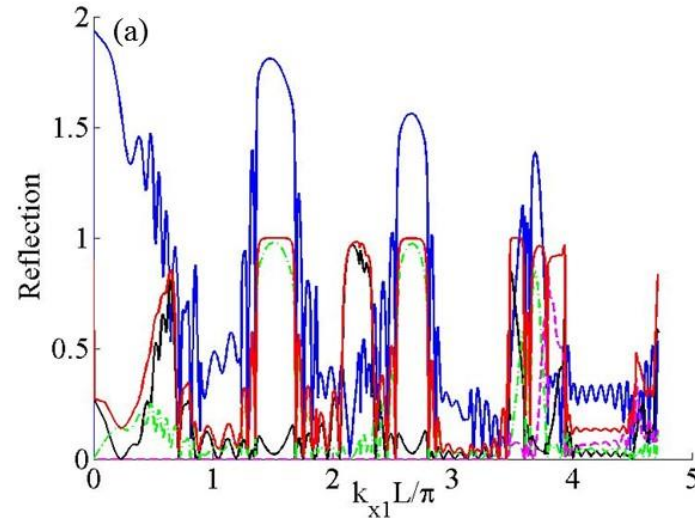
Reflection coefficient in the case of sparse array of 11 cylinders ( $a/L = 0.10$ ,  $L/W = 1$ ) versus  $kL/\pi$ : (- -), 1st mode,  $P = 2$  (3 modes); (-), 1st mode,  $P = 3$  (4 modes); (- -), second mode,  $P = 2$  (3 modes), (-), second mode,  $P = 3$  (4 modes); (- -), third mode,  $P = 2$  (3 modes); (-), third mode,  $P = 3$  (4 modes).

# Propagation in the presence of periodic cylinders :

## Bragg resonances for scattered waves

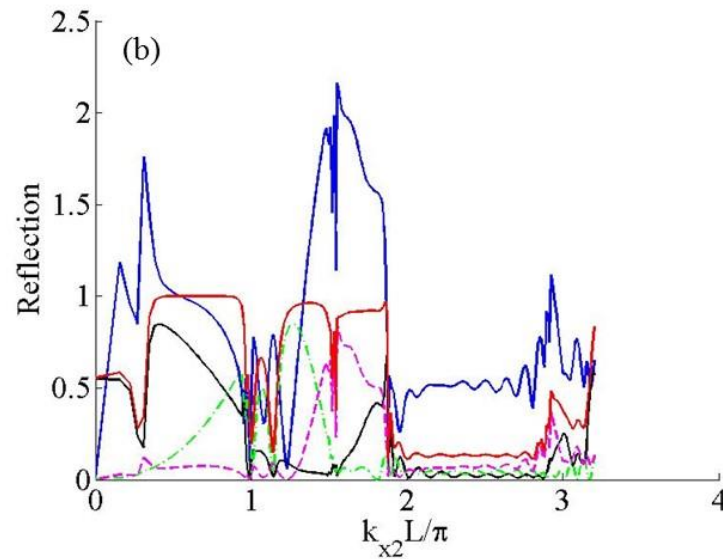
(a) Reflection versus  $k_{x1}L/\pi$ :

- (-), reflection coefficient for the second mode;
- (-), total reflected energy;
- (-), reflected energy for the 1st mode;
- (-.-), reflected energy for the second mode;
- (-.-), reflected energy for the third mode;



(b) Reflection versus  $k_{x2}L/\pi$ :

- (-), reflection coefficient for the third mode;
- (-), total reflected energy;
- (-), reflected energy for the 1st mode;
- (-.-), reflected energy for the second mode;
- (-.-), reflected energy for the third mode.



Resonance conditions:

$$k_{xn}L = (2p + 1)\pi/2, \quad p > 1 \text{ where } n = 1 \text{ for mode 2 and } n = 2 \text{ for mode 3}$$

# Application to wave energy device : Oscillating water column (OWC)

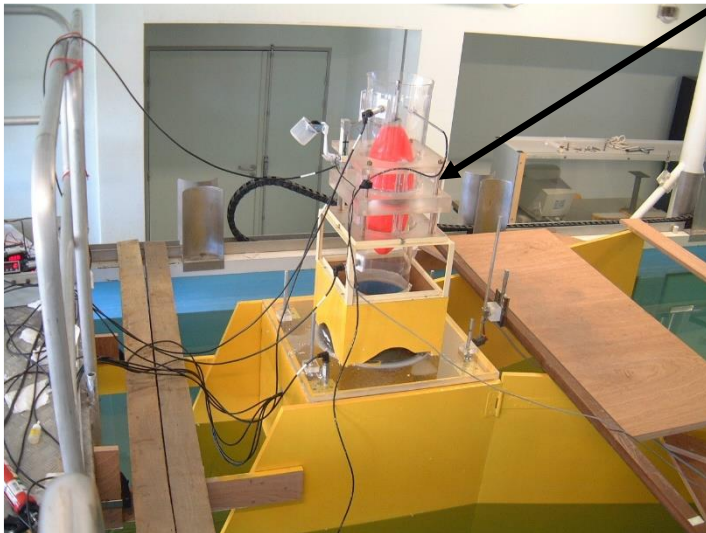
Research project at ISITV, UTLN (years 2000):

How to provide energy, for punctual needs in the natural park of Scandola in Corsica, France for public lighting?

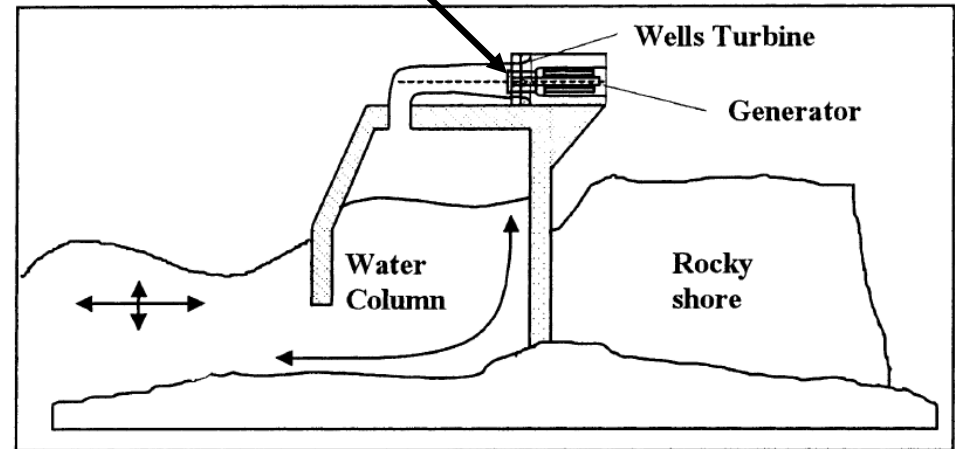
Bay of Girolata, Corsica, France



Wells turbine



Prototype, ISITV (now SeaTech), UTLN

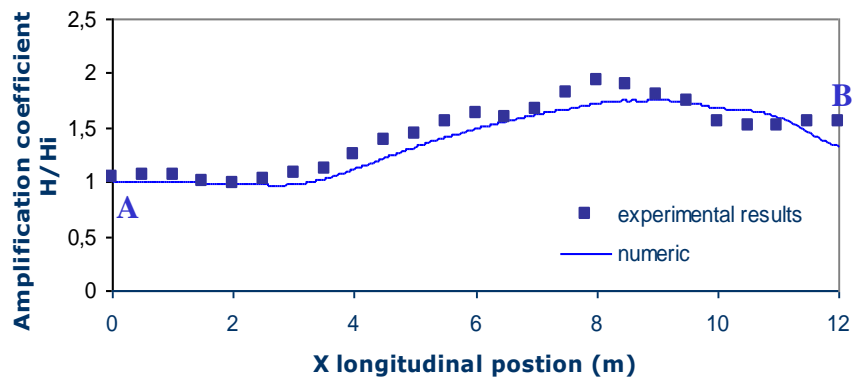
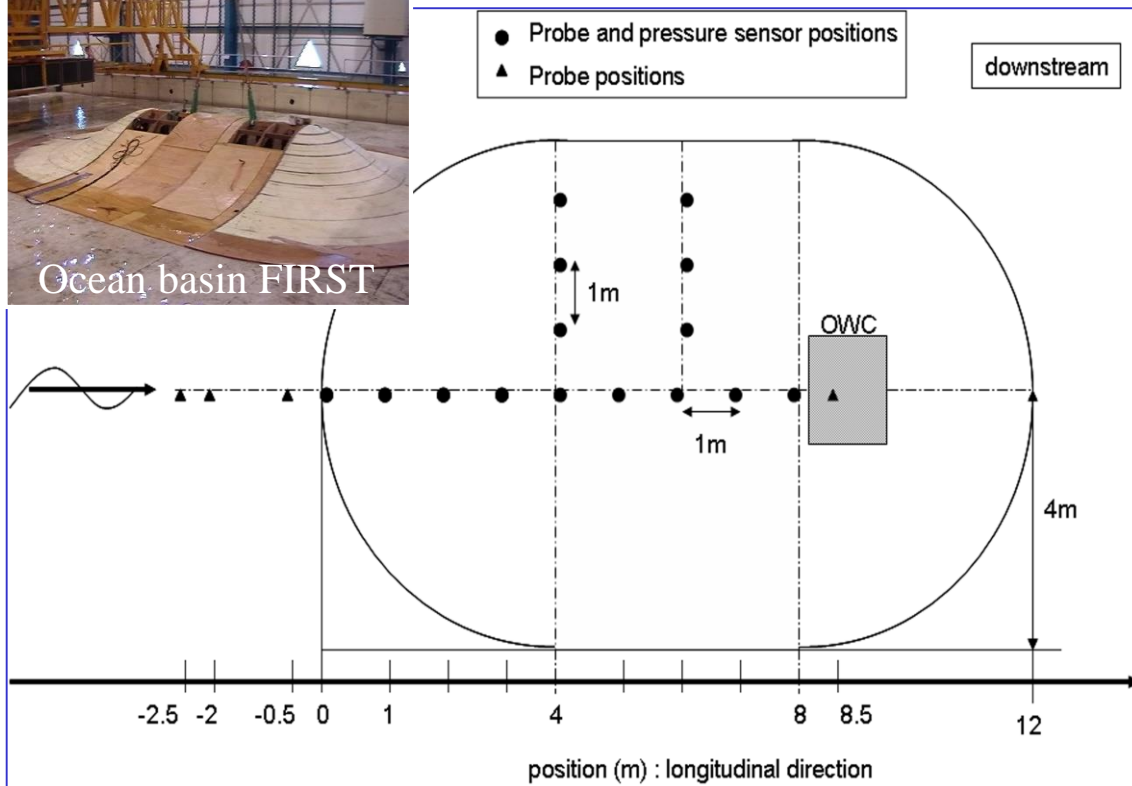


How to increase incoming wave energy?

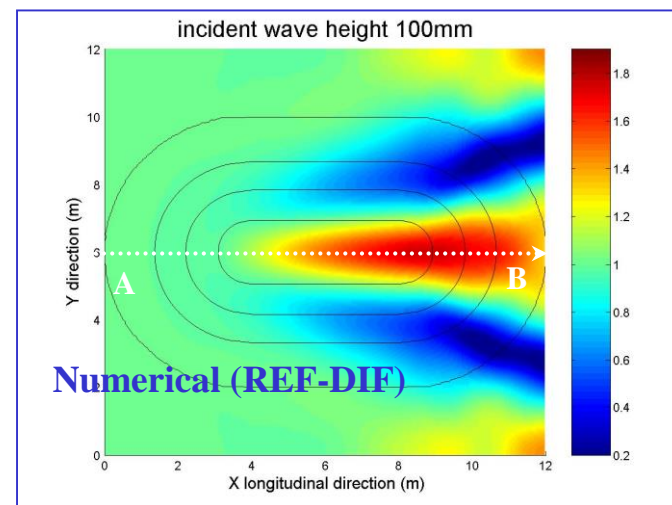
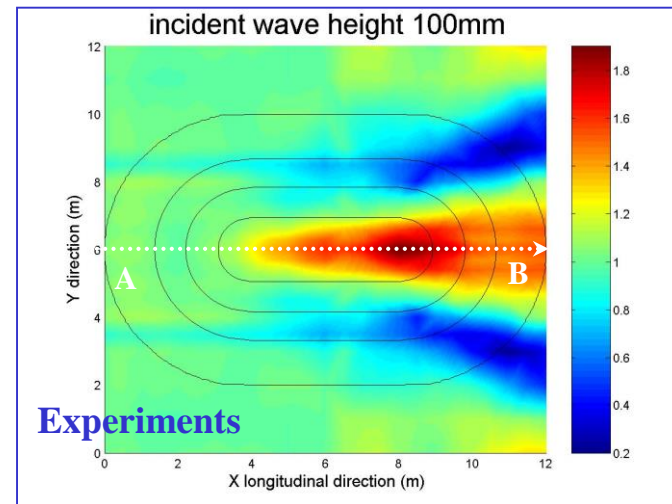
➔ Wave focusing above a shoal

Oscillating water column  
(from Delauré and Lewis, 2003)

# Application to wave energy device : Oscillating water column (OWC)



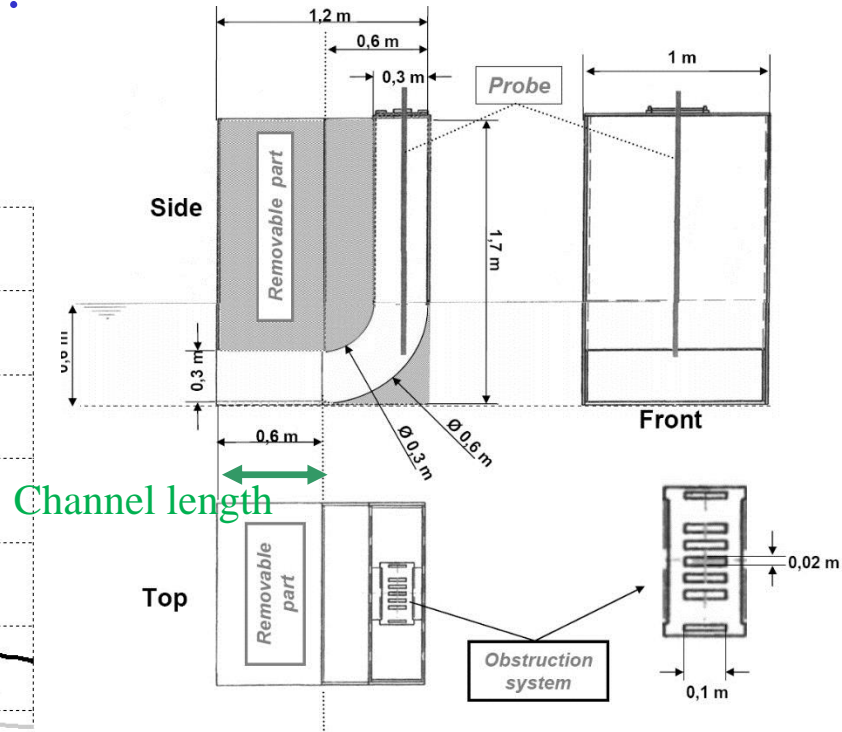
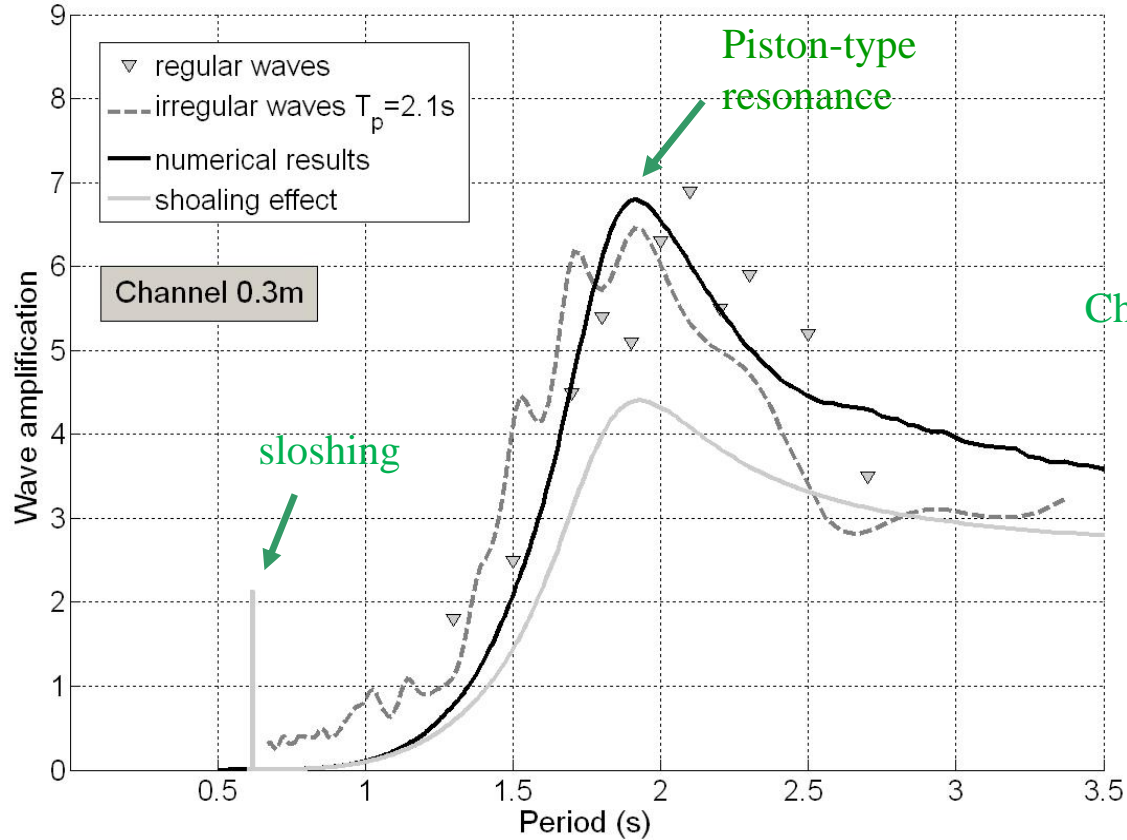
$T=1.7s$



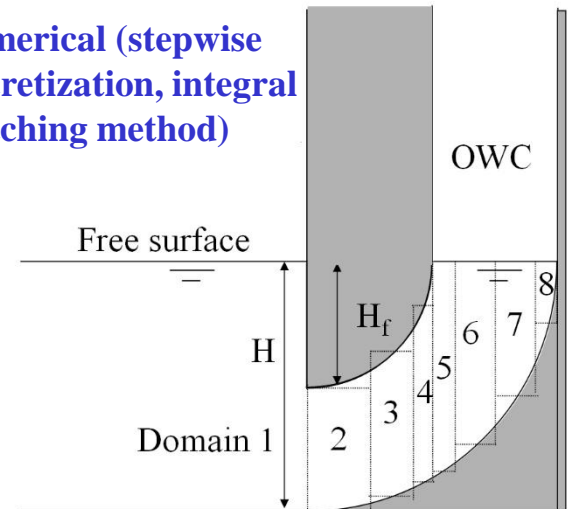


# Application to wave energy device :

## Wave amplification



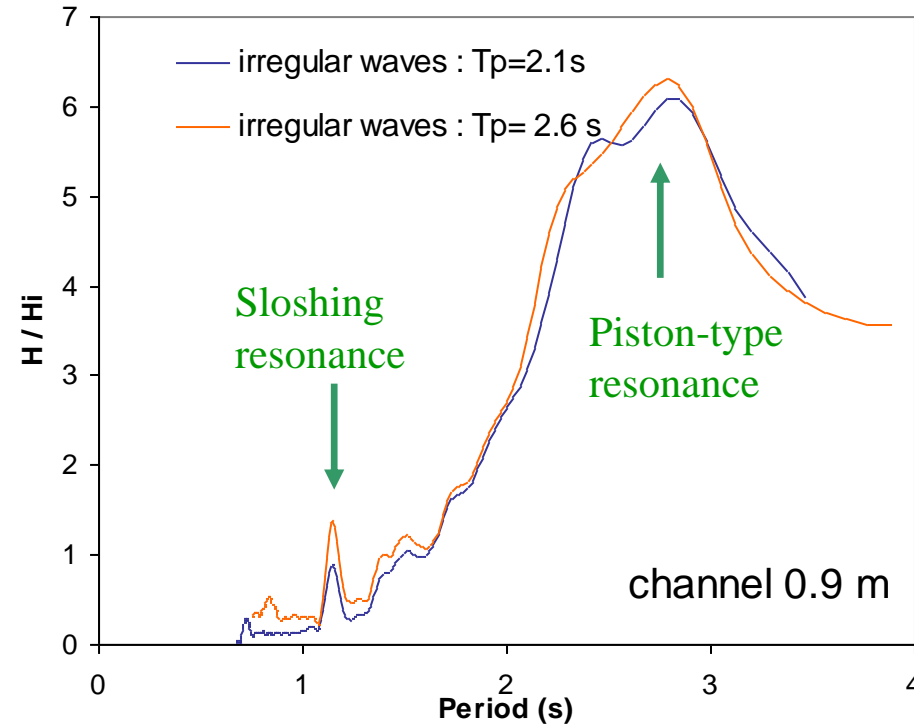
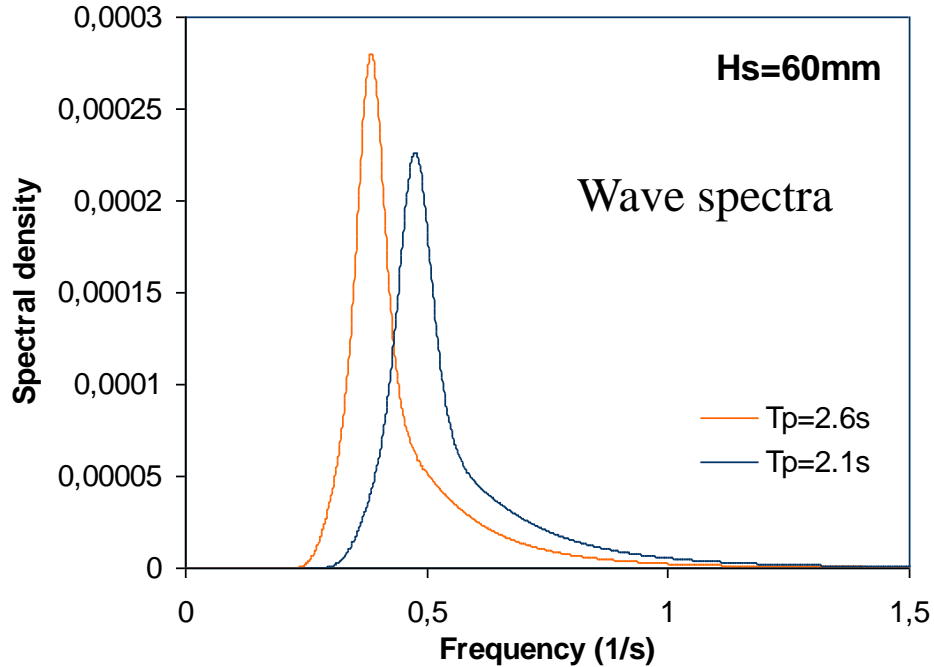
**Numerical (stepwise discretization, integral matching method)**



- ➡ Wave amplification depends on the geometry of the OWC
- ➡ Wave amplification enhanced by the shoal

# Wave amplification in the water column

## Linear wave behaviour



- Use of two Jonswap spectra
- Both transfer functions are almost identical

Non-linear effects are rather negligible in that case

# Conclusions

- **Inhomogeneous medium : water wave celerity changes**  
(bathymetry, current, porous medium, presence of surface-piercing structures)
- **Scattering :**  
2D, reflection, transmission  
3D, + refraction, diffraction, focusing or dispersion  
**Leads to**  
**Interference process** due to wave partial reflection at domain boundaries  
and **Bragg resonance** for periodic inhomogeneous media
- **Irregular waves :**  
interference process less important due to  
wave energy spreading on both frequencies and direction of propagation
- In the 3d Case of constant water depth :  
**Helmotz equation**, **analogy to other types of waves as acoustic waves**

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